



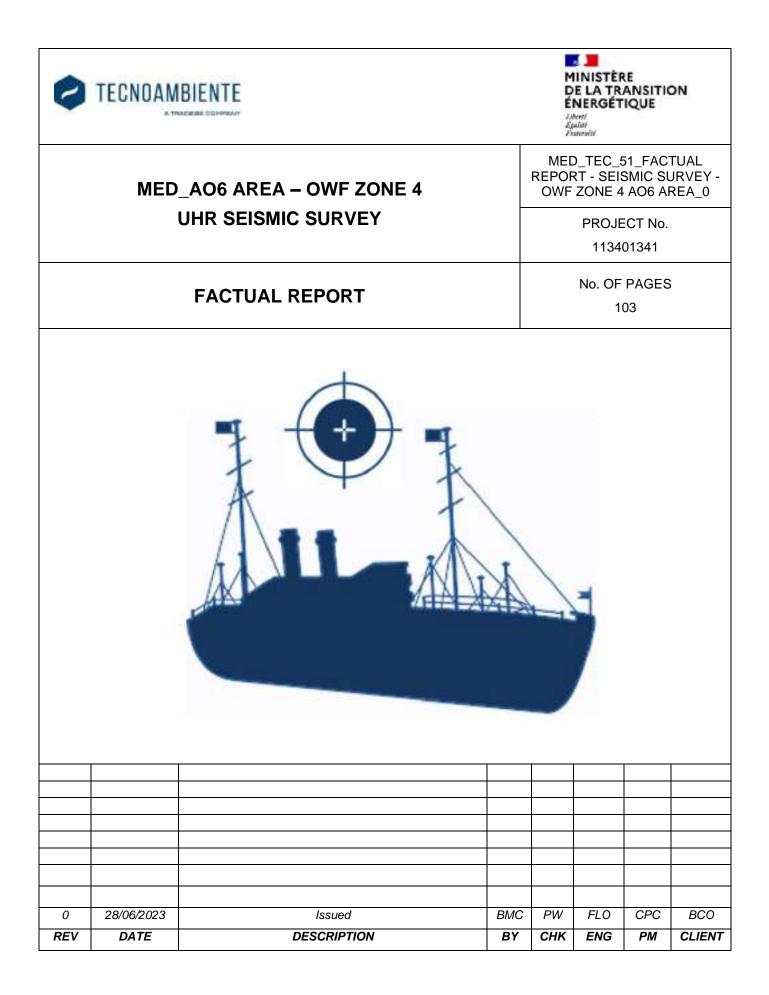
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MINISTÈRE DE LA TRANSITION	Project	Package	Issuer	Chrono	Revision	Status	
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	Title	MED_TEC_51_Factual report - Seismic survey - OWF Zone 4 AO6 area_0					

CONTENTS

EX	ECUTI\	/E SUMMARY	6
1.	INTR	ODUCTION	6
1	.1.	PROJECT OVERVIEW	6
1	.2.	SCOPE OF WORK	10
1	.3.	GEODETIC PARAMETERS	10
	1.3.1.	Survey datum	10
	1.3.2.	Vertical datum	11
	1.3.3.	Tidal reduction	11
2.	METH	HODOLOGY	13
2	2.1.	MBES BATHYMETRY	13
	2.1.1.	Data acquisition	13
	2.1.2.	Data processing	15
2	2.1.	UHR SEISMIC	16
	2.1.1.	Data acquisition	16
	2.1.2.	Data processing	19
3.	RESU	JLTS	23
3	8.1.	BATHYMETRY	23
Э	8.2.	GEOLOGY	27
	3.2.1.	Geological setting from background data	27
	3.2.2.	Geological sequence	27
	3.2.3.	Geohazards	39
	3.2.4.	Background data summary and regional geology	41
	3.2.5.	Conclusions and recommendations/comments	44

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 1 of 49
Issue date	28/06/2023	Document uncontrolled when		trolled when printed/downloaded

MINISTÈRE DE LA TRANSITION	Project	Package	Issuer	Chrono	Revision	Status	
	MED_AO6	GPY	TEC	51	0	IFR	
	Title	MED_TEC_51_Factual report - Seismic survey - OWF Zone 4 AO6 area_0					

4.	REFERENCES	45
APP	PENDIX I – CHARTING	46
APP	PENDIX II – SEISMIC PROCESSING OVERVIEW	47

LIST OF FIGURES

Figure 1-1: MED_AO6 survey area7
Figure 1-2: Windfarm area (OWF) in the MED_AO6 Zone 4 Survey area8
Figure 1-3: Line plan for MED_AO6 Zone 4 Windfarm area (OWF)9
Figure 1-4: QINSy's method for accurate tide calculation12
Figure 2-1: MBES bathymetry data acquisition with the QINSy software14
Figure 2-2: Processing screen of MBES bathymetry data with the Qimera software15
Figure 2-3: 3D image of the MBES bathymetry processing15
Figure 2-4: MBES bathymetry processing overview16
Figure 2-5: Screenshot from Geometrics LH16 software during UHR Seismic acquisition.
Figure 2-6: UHR Seismic processing overview20
Figure 13-7: Minimum phase brute stack21
Figure 13-8: Zero-phase corrected control stack (Demultiple, Noise filtering, Deghost, Static
Correction, Far Trace Mute & Pre-Stack Migration)22
Figure 13-9: Zero-phase corrected final processed stack (Time-Variant Bandpass filter, FK
filter, Gain balancing)22
Figure 3-1: Colour table for the representation of the MBES terrain model23
Figure 3-2: Whole bathymetric data grid model (1 x 1 m) for the MED_AO6 Zone 4 OWF.
Figure 3-3: Detail of the bathymetric data grid model (1 x 1 m) for the offshore windfarm of
the MED_AO6 Z3 area – Survey lines and tracklines MED_OWF_Z3_002-03525
Figure 3-4: Detail of the bathymetric data grid model (1 x 1 m) for the offshore windfarm of
the MED_AO6 Z3 area – Survey lines and tracklines MED_OWF_Z3_022-02725

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 2 of 49
Issue date	28/06/2023	Document uncontrolled when printed/downloaded		trolled when printed/downloaded

MINISTÈRE DE LA TRANSITION	Project	Package	Issuer	Chrono	Revision	Status	
	MED_AO6	GPY	TEC	51	0	IFR	
	Title	MED_TEC_51_Factual report - Seismic survey - OWF Zone 4 AO6 area_0					

Figure 3-5: Detail of the bathymetric data grid model (1 x 1 m) for the offshore windfarm of
the MED_AO6 Z3 area – Survey lines and tracklines MED_OWF_Z3_006-00726
Figure 3-6: Detail of the bathymetric data grid model (1 x 1 m) for the offshore windfarm of
the MED_AO6 Z3 area – Survey lines and tracklines MED_OWF_Z3_022-00726
Figure 3-7: Line AO6_OWF_Z4_018
Figure 3-8: Line AO6_OWF_Z4_00732
Figure 3-9: Isopach of depth to Base Unit 1 (H05)34
Figure 3-10: Isopach of depth to Base Unit 2 (H10)35
Figure 3-11: Isopach of depth to Base Unit 3 (H20)36
Figure 3-12: Isopach of depth to Base Unit 4 (H30)37
Figure 3-13: Large channel extents within Zone 4 OWF area
Figure 3-14: Figure showing the geology (Figure 3 from Rabineau et al., 2003;
Benabdellouahed, 2011; Paquet et al b, in preparation)43

LIST OF TABLES

Table 1: Datum parameters table	10
Table 2: Projection parameters table	11
Table 3: UHRS operational parameters	17
Table 4: Shallow geological units.	
Table 5: Geological characteristics / processes and potential constraints	

ABBREVIATIONS

ADCP	Acoustic Doppler Current Profiler
ASV	Assumed Sound Velocity
Ch	Channel
cm	Centimetre
СМР	Common Mid-Point
C-0	Computed Minus Observed
CoG	Centre of Gravity
CRP	Central Reference Point

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 3 of 49
Issue date	28/06/2023	Docun	nent uncon	trolled when printed/downloaded

Project	Package	Issuer	Chrono	Revision	Status	
MED_AO6	GPY	TEC	51	0	IFR	
Title	MED_TEC_51_Factual report - Seismic survey - OWF Zone 4 AO6 and					

ABBREVIATIONS

CSP	Central Shot Point
DEMOB	Demobilisation
DGEC	Direction générale de l'énergie et du climat
DP	Dynamic Positioning
DPO	Dynamic Positioning Officer
DPR	Daily production report
EP	Environmental Protection
FLO	Fisheries Liaison Officer
GNSS	Global Navigation Satellite System
GPS	Global Positioning System
GRS	Geodetic Reference System
GSO	Geophysical Services Offshore
h	Hour
IMO	International Maritime Organization
J	Joule
JNCC	Joint Nature Conservation Committee
kHz	Kilohertz
LAT	Low Astronomical Tide
m	Meters
min	Minutes
MBES	Multibeam echosounder
mm	Millimetre
МОВ	Mobilisation
MRU	Motion Reference Unit
РОВ	Personnel On Board
PAM	Passive Acoustic Monitoring
PPP	Precise Point Positioning
PPSU	Pulse Power Supply Unit
QA-QC	Quality Assurance – Quality Control

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 4 of 49
Issue date	28/06/2023	Docun	nent uncon	trolled when printed/downloaded

Project	Package	Issuer	Chrono	Revision	Status	
MED_AO6	GPY	TEC	51	0	IFR	
Title	MED_TEC_51_Factual report - Seismic survey - OWF Zone 4 AO6 area					

ABBREVIATIONS

RTE	Réseau de Transport d'Électricité
RTK	Real Time Kinematics
S	Second
SHOM	Service Hydrographique et Océanographique de la Marine
SN	Serial Number
SRF	Ship's Reference Frame
SVP	Sound Velocity Profiler
SVS	Sound Velocity Sensor
ТВС	To be confirmed
TTS	TTSurvey Ltd (Seismic equipment hire company)
UHR	Ultra-High Resolution
UTC	Coordinated Universal Time
UTM	Universal Transverse Mercator
VSAT	Very-Small-Aperture Terminal
WB	Water Bottom
WD	Water Depth
WGS84	World Geodetic System 1984
WT	Work time
ZH	Hydrographic Zero or Hydrographic Datum

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 5 of 49
Issue date	28/06/2023	Docun	nent uncon	trolled when printed/downloaded

MINISTÈRE DE LA TRANSITION	Project	Package	Issuer	Chrono	Revision	Status
	MED_AO6	GPY	TEC	51	0	IFR
	Title MED_TEC_51_Factual report - Seismic survey - OWF Zone 4 AO6 area					

EXECUTIVE SUMMARY

According to the detected bathymetric results in the acquired lines of the offshore windfarm survey area, the depth range correspond between -86.78 meters in the shallowest part at the north-eastern region, to -112.85 meters in the deepest part at the southern region.

The seabed slopes are gentle, with average seabed gradients of 1.07°, maximum values of 28.17° and minimum values of 0°. The seabed gradient was computed at 22 planned borehole locations distributed over the range of water depths across the site.

Late Pleistocene sediments, comprising interbeds of soft silty clay and coarser reworked sands occur throughout the OWF area. Unconsolidated sediments extend to deeper than the depth of interest within the OWF survey area (60m+) in thickness, and the units all dip gently towards the southeast.

No shallow gas, or other geohazards are expected within the OWF area.

1. INTRODUCTION

1.1. PROJECT OVERVIEW

Tecnoambiente carried out four geophysical surveys within the proposed MED_AO6 lot located along the southeast coast of the French Mediterranean shore in the Gulf of Lion. The areas of interest are located in the Gulf of Lion off the French Mediterranean coast. These areas are 4 offshore windfarm (Zone 1 OWF, Zone 2 OWF, Zone 3 OWF and Zone 4 OWF) and 3 offshore substations (Zone 1 OSS, Zone 2 OSS and Zone 3 OSS) which are under investigation in this project (Figure 1-1). Each site is under consideration for a windfarm and offshore substation.

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 6 of 49
Issue date	28/06/2023	Docun	nent uncon	trolled when printed/downloaded

MINISTÈRE DE LA TRANSITION	Project	Package	Issuer	Chrono	Revision	Status
	MED_AO6	GPY	TEC	51	0	IFR
						4 AO6 area_0

The area of relevance in this report is Zone 4, located off the coast of Marseille (Figure 1-2):

- Area: 267.35 km²
- *Dimensions:* 21.40 km x 20.39 km.
- Bathymetric range: -87 m to -113 m (Vertical reference Bathyelli v2 ZH)

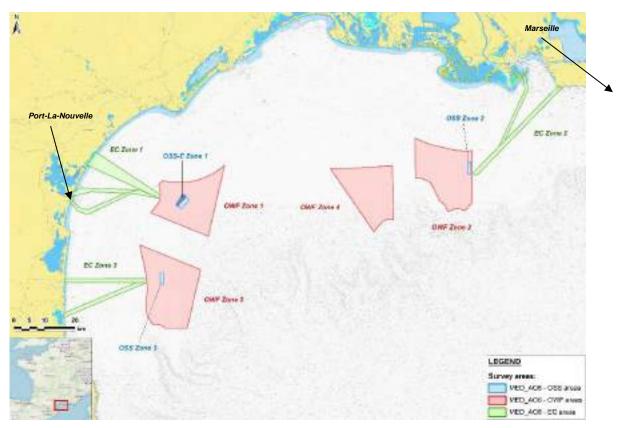


Figure 1-1: MED_AO6 survey area.

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 7 of 49
Issue date	28/06/2023	Document uncontrolled when printed/downloaded		trolled when printed/downloaded

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	MED_AO6	GPY	TEC	51	0	IFR	
	Title	MED_TEC_51_Factual report - Seismic survey - OWF Zone 4 AO6 area_0					

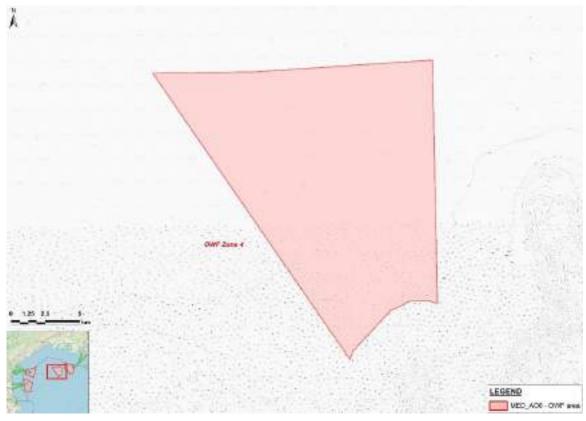


Figure 1-2: Windfarm area (OWF) in the MED_AO6 Zone 4 Survey area.

The following data were used in the study:

• 392.40 km of MBES and UHRS data

The goals of this first phase of the geophysical surveys are to perform

- 1. 2D UHR surveys in the wind farm areas in zones 1 to zone 4,
- 2. 2D UHR survey on offshore substations in zones 1 to zone 3.

Figure 1-3 shows the survey line plan.

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 8 of 49
Issue date	28/06/2023	Document uncontrolled when printed/downloaded		trolled when printed/downloaded

	Project	Package	Issuer	Chrono	Revision	Status
	MED_AO6	GPY	TEC	51	0	IFR
	Title	MED_TEC_51_Factual report - Seismic survey - OWF Zone 4 AO6 area_0				

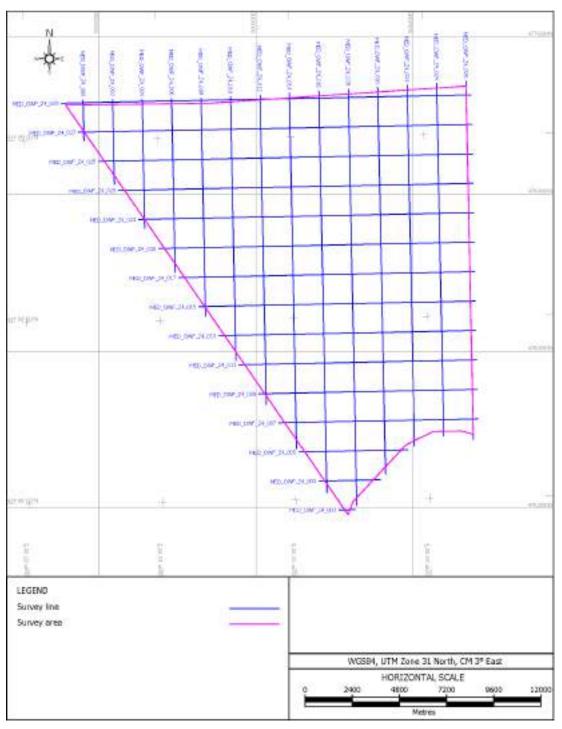


Figure 1-3: Line plan for MED_AO6 Zone 4 Windfarm area (OWF).

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 9 of 49
Issue date	28/06/2023	Docum	nent uncon	trolled when printed/downloaded

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	MED_AO6	GPY	TEC	51	0	IFR	
	Title	MED_TEC_51_Factual report - Seismic survey - OWF Zone 4 AO6 area_0					

1.2. SCOPE OF WORK

The objective of the site survey was to perform a geophysical survey over the proposed OWF and OSS sites comprising MBES, SBP and UHRS datasets. The purpose of this was to:

- To define the water depths and seabed topography.
- To define the shallow (nominally 30 m BSB) subsurface geology.
- Review proposed borehole locations for geohazards.

The main purpose of the study is to provide an interpretation of the geophysical data to better understand the main characteristics of the seabed and geology at the project location and to undertake a derisking study over the MED_AO6 Zone 4 OWF site.

1.3. GEODETIC PARAMETERS

1.3.1. Survey datum

These parameters are detailed below.

Table 1: Datum parameters table

DATUM	
Survey Datum:	WGS 84
Spheroid	GRS 1980
Semi-Major Axis (a)	6378137.00000000
Semi-Minor Axis (b)	6356752.314245179
Inverse Flattening (1/f)	1/298.257223563

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 10 of 49
Issue date	28/06/2023	Document uncontrolled when printed/downloaded		trolled when printed/downloaded

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	MED_AO6	GPY	TEC	51	0	IFR
	Title	MED_TEC_51_Factual report - Seismic survey - OWF Zone 4 AO6 area_0				

Table 2: Projection parameters table.

PROJECTION				
Projection	UTM			
False Easting	500000			
False Northing	0			
Latitude of Origin	0°00'00.00000''			
Central Meridian	3°00'00.000000''			
UTM Zone	31 N			
Scale Factor on CM	0.9996			
Units:	Meters			

1.3.2. Vertical datum

The vertical datum used in the QINSy software is Bathyelli v2.0 ZH geoid published by the SHOM in December 2013. The Bathyelli v2.0 ZH (SHOM 2013) is a surface based on the GRS 1980 spheroid, and it is a set of surfaces each of which defines the separation of one vertical datum from the WGS84 ellipsoid to the vertical maritime reference Hydrographic Datum or Hydrographic Zero. These ellipsoidal heights are given in meters.

This geoid covers the intersection between the SHOM tidal model and the different tidal zones of France.

For the survey area MED_AO6 Z4, the corrections to hydrographic zero are made by tidal observations of the port of Marseille (Corniche) ($43^{\circ}16'$ N – 05° 21' E). The difference between the hydrographic zero and the LAT reference level for this port is 0.27 m, according to the study by SHOM "*Références Altimétriques Maritimes. Ports de France métropolitaine et d'outre-mer*" of 2019.

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 11 of 49
Issue date	28/06/2023	Document uncontrolled when printed/downloaded		trolled when printed/downloaded

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	MED_AO6	GPY	TEC	51	0	IFR
	Title	MED_TEC_51_Factual report - Seismic survey - OWF Zone 4 AO6 area_0				

1.3.3. Tidal reduction

In order to carry out the survey as accurately as possible, Tecnoambiente was receiving MarineStar PPP corrections by satellite signal. When using an accurate GNSS system the tidal corrections were carried out in real-time through QINSy computations, as it is shown in the next figure.

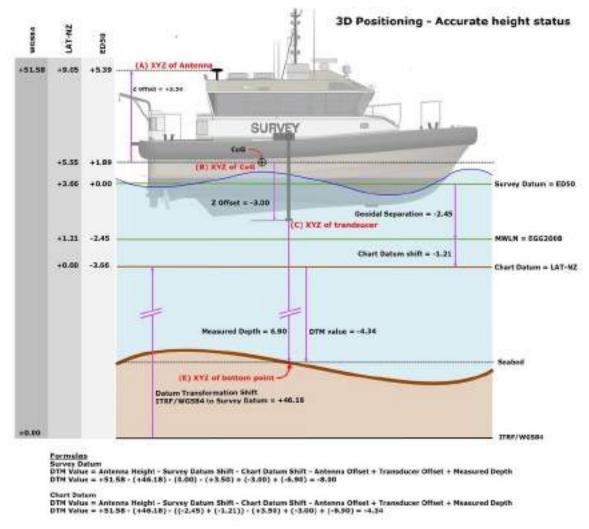


Figure 1-4: QINSy's method for accurate tide calculation.

In the event that corrections drop out they can be applied in post processing.

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 12 of 49
Issue date	28/06/2023	Docum	nent uncont	trolled when printed/downloaded

MINISTÈRE DE LA TRANSITION	Project	Package	Issuer	Chrono	Revision	Status
	MED_AO6	GPY	TEC	51	0	IFR
Title MED_TEC_51_Factual report - Seismic survey - OWF Zone 4 A					4 AO6 area_0	

2. METHODOLOGY

2.1. MBES BATHYMETRY

2.1.1. Data acquisition

The objective during the data acquisition is the referencing of the acquired seismic data, therefore, the total coverage of the study area was not necessary. Due to this the project lines have been designed with a spacing of 1500 meters.

During the data acquisition, the vessel's master must follow the previously programmed routes of the project lines, governed by the indications of the computer screen (Helmsmann indicator), which is shown, by means of visual and audible alarms, when it separates from its course more than a specified amount (variable according to weather conditions in the area, but never more than 2.5 metres from the theoretical line), and also when there is a problem in a peripheral, such as the loss of GPS corrections.

While the master follows the navigation lines, the acquisition module of the hydrography program captures all the position data sent by the GPS, as well as the soundings sent by the multibeam sounder for each transmission pulse, as well as the values of the heading, wave height, roll and head angles sent by the MRU.

Parallel to the data entry, the data acquired by the equipment and peripherals is synchronized. This process is carried out by QINSy itself, complemented by the input of the time and the pulse per second (PPS) provided by the MRU, so that all the data is time synchronised.

The guidelines followed by Tecnoambiente during the survey for MBES data acquisition were the following ones:

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 13 of 49
Issue date	28/06/2023	Docun	nent uncont	trolled when printed/downloaded

MINISTÈRE DE LA TRANSITION	Project	Package	Issuer	Chrono	Revision	Status
	MED_AO6	GPY	TEC	51	0	IFR
Title MED_TEC_51_Factual report - Seismic survey - OWF Zor					4 AO6 area_0	

- IT-CM-01. Guidelines for Hydrography Project management, 5
- IT-CM-04. Bathymetric survey, 1
- IT-CM-14 Survey Basics Guidance, 1
- IT-CM-15 Online Surveying procedure, 3

These guidelines can be found in the quality plan document MED_TEC_04_Quality Plan.

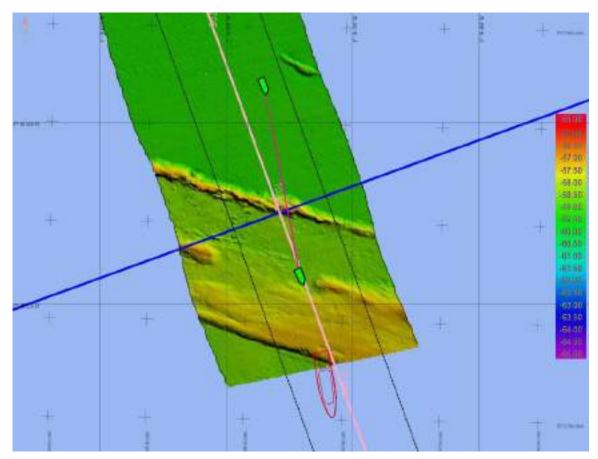


Figure 2-1: MBES bathymetry data acquisition with the QINSy software.

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 14 of 49
Issue date	28/06/2023	Document ur		trolled when printed/downloaded

MINISTÈRE DE LA TRANSITION	Project	Package	Issuer	Chrono	Revision	Status	
	MED_AO6	GPY	TEC	51	0	IFR	
	Title	MED_TEC_51_Factual report - Seismic survey - OWF Zone 4 AO6 area_0					

2.1.2. Data processing

A single head Kongsberg EM 2040 high resolution MBES system that is permanently installed on the Geo Focus vessel was used to produce digital terrain models (DTMs).

Along the processing phase of the acquired data, the lines on the screen are processed in order to manually match the height of the bathymetric lines and also correct the noise that appears in the records, noise produced by multiple factors such as, multipath in position, air bubbles, motor interference of the vessel etc. in the digital register of soundings.

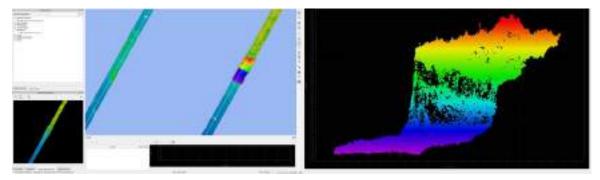


Figure 2-2: Processing screen of MBES bathymetry data with the Qimera software.

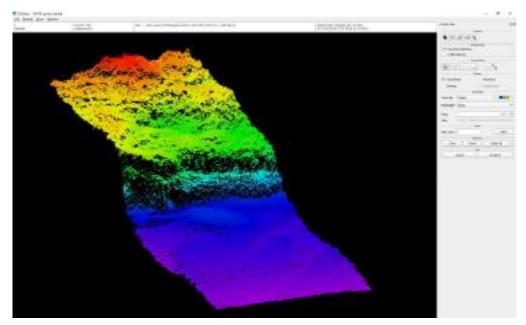


Figure 2-3: 3D image of the MBES bathymetry processing.

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 15 of 49
Issue date	28/06/2023	Docun	nent uncon	trolled when printed/downloaded

MINISTÈRE DE LA TRANSITION	Project	Package	Issuer	Chrono	Revision	Status
ÉNERGÉTIQUE Uter Aplan Franshi	MED_AO6	GPY	TEC	51	0	IFR
Title MED_TEC_51_Factual report - Seismic surve					survey - OWF Zone	4 AO6 area_0

Once the possible existing errors in the records have been deleted, a digital model of the terrain with 1×1 m grid size has been made with a minimum cell size to obtain the maximum resolution of the seabed.

The general MBES processing workflow is presented in the following figure.

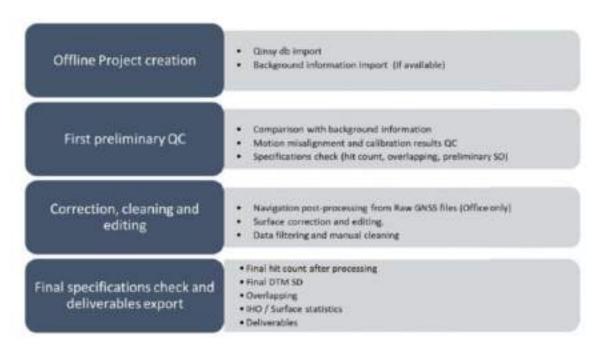


Figure 2-4: MBES bathymetry processing overview.

2.1. UHR SEISMIC

2.1.1. Data acquisition

UHR Seismic data was acquired using GSO 400-tip Sparker sled and Applied Acoustics CSP-N pulsed power supply unit were mobilised as the acquisition source, interfaced with a Geometrics GEOEEL LH16 recording system and 48 channel UHR streamer. The first 24 channels of the streamer at 1m group interval and the remaining 24 channels at 2m. The streamer was kept at a depth of 1m by a head and tail buoy as well as 2 Digicourse 5011 levellers (Birds).

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 16 of 49
Issue date	28/06/2023	Docun	nent uncont	trolled when printed/downloaded

MINISTÈRE DE LA TRANSITION	Project	Package	Issuer	Chrono	Revision	Status	
	MED_AO6	GPY	TEC	51	0	IFR	
	Title	MED_TEC_51_Factual report - Seismic survey - OWF Zone 4 AO6 area_0					

Accurate positioning was collected using Modulus 101G GPS pods mounted on each towed system, Sparker sled as well as head & tail buoys for streamer positioning.

The shot point interval for the survey was 1 m, giving a nominal fold of 36 when binning with a CDP spacing of 1m to keep the bins consistent with the variable channel spacing. True fold will vary around this value when real source and receiver positions are used rather than nominal geometry, according to variations in ship speed and feather angle changes between shots.

Parameter	Value		
Active Streamer Length	75m		
Number of channels	48		
Group Length	Channels 1-24: 1 m		
	Channels 25-48: 2 m		
Target Tow Depth	1m +/-0.5m		
Near Offset	~5-6m		
Sample Rate	0.0625ms		
Record Length	0.250ms		
Shot Point Interval	1m		
Source	Sparker – GSO – 400 tips		
Target Source Tow Depth	0.3 m		

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 17 of 49
Issue date	28/06/2023	Docun	nent uncon	trolled when printed/downloaded

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	Title	MED_TEC_5	1_Factual re	port - Seismic	survey - OWF Zone	4 AO6 area_0

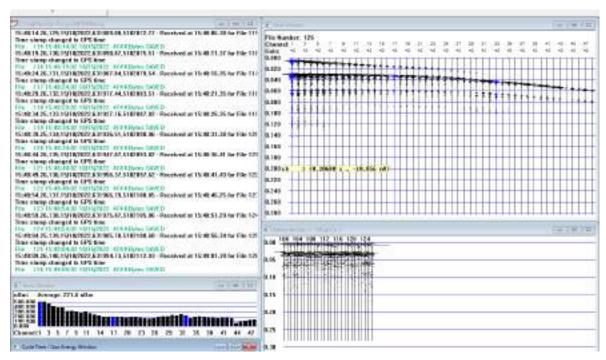


Figure 2-5: Screenshot from Geometrics LH16 software during UHR Seismic acquisition.

The guidelines followed by Tecnoambiente during the surveying for UHR Seismic data acquisition were the ones provided from TTS, which are:

- IT-CM-17 Sparker Deployment Recovery, 1
- IT-TTS- 01_Geoeel Instrument Verification Procedure
- IT-TTS-02_Multichannel Seismic Streamer Procedure
- IT-TTS-03_Sparker Pulse Test Procedures
- IT-TTS-04_Streamer Recovery

These guidelines can be found in the quality plan document MED_TEC_04_Quality Plan.

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 18 of 49
Issue date	28/06/2023	Docun	nent uncon	trolled when printed/downloaded

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	MED_AO6	GPY	TEC	51	0	IFR		
	Title	MED_TEC_51_Factual report - Seismic survey - OWF Zone 4 AO6 area_0						

2.1.2. Data processing

Data processing and interpretation was performed within the MED_AO6 area to a recorded length of 100 m sub-seabed for both OWFs and OSSs. This interpretation was done for evaluation of seabed and sub-seabed conditions.

The dataset was quality controlled offshore on board the vessel Geo Focus by Peak Processing using a Linux based system with Landmark's ProMAX/SeisSpace processing software.

The dataset was then made available to Peak Processing upon completion of the fieldwork with the processing of the raw UHR seismic data performed and finalised using Shearwater Reveal version 5.1 on a small cluster.

Stacking velocities generated during the processing of the UHR data were used to help choose velocities in the time-depth calculation. Standardised velocities were chosen based on the sediment characteristics expected. For the interbedded sands and clays, as interpreted in AO6, an assumed seismic velocity of 1700m/s has been used.

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 19 of 49
Issue date	28/06/2023	Docun	nent uncon	trolled when printed/downloaded

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	MED_AO6	GPY	TEC	51	0	IFR
	Title	MED_TEC_5	1_Factual re	port - Seismic	survey - OWF Zone	4 AO6 area_0

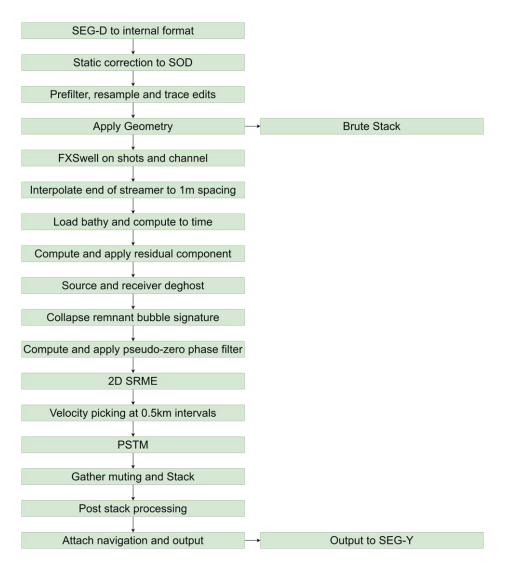


Figure 2-6: UHR Seismic processing overview.

For quality assurance, displays of the following were produced for each line with a copy provided to the client representative offshore, in addition to the Brute Stack SGY exported:

- Near trace
- Shot record examples (displayed every 100 shots)
- RMS Noise Display (calculated every 100 shots)
- Spectral Analysis

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 20 of 49
Issue date	28/06/2023	Docun	nent uncon	trolled when printed/downloaded

	Project	Package	Issuer	Chrono	Revision	Status
	MED_AO6	GPY	TEC	51	0	IFR
	Title	MED_TEC_5	1_Factual re	port - Seismic	survey - OWF Zone	4 AO6 area_0

- Offset QC checks, showing computed arrival time from offsets derived from GPS navigation data overlaid on top of the direct arrival in the data itself
- Velocity Semblance/Gather Example
- Brute Stack, annotated with trace fold header plot

Processing of the dataset took place ashore, and full details of the processing can be found in a dedicated seismic processing report. As a summary, the evolution of the dataset can be seen in the following examples shown at three different stages of the processing:

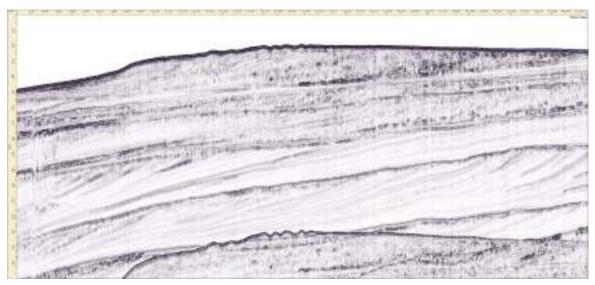


Figure 2-7: Minimum phase brute stack.

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 21 of 49
Issue date	28/06/2023	Docun	nent uncon	trolled when printed/downloaded

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	MED_AO6	0	IFR			
	Title	MED_TEC_51_Factual report - Seismic survey - OWF Zone 4 AO6 area_0				

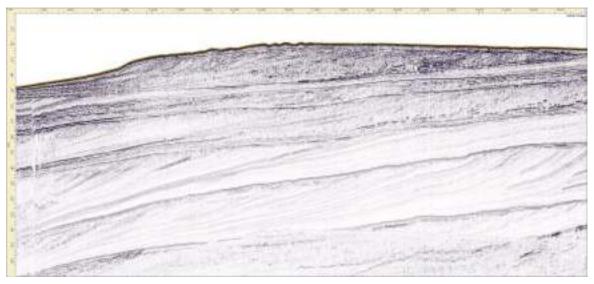


Figure 2-8: Zero-phase corrected control stack (Demultiple, Noise filtering, Deghost, Static Correction, Far Trace Mute & Pre-Stack Migration).

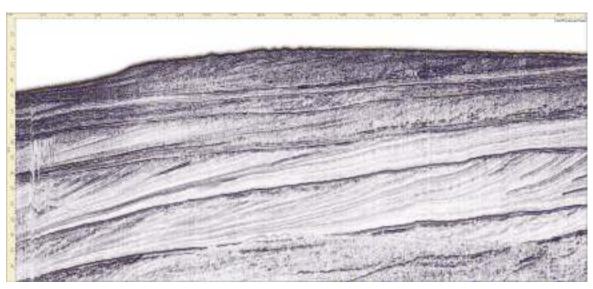


Figure 2-9: Zero-phase corrected final processed stack (Time-Variant Bandpass filter, FK filter, Gain balancing).

Interpretation was cross-checked for consistency at all crossline locations. Additionally, geohazards assessment was carried out focusing on the areas for planned shallow geotechnical operations. Heave corrections are applied after datum alignment with the MBES data.

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 22 of 49
Issue date	28/06/2023	Docun	nent uncont	trolled when printed/downloaded

MINISTÈRE DE LA TRANSITION	Project	Package	Issuer	Chrono	Revision	Status		
	MED_AO6	GPY	TEC	51	0	IFR		
	Title	MED_TEC_51_Factual report - Seismic survey - OWF Zone 4 AO6 area_0						

3. RESULTS

3.1. BATHYMETRY

The detailed resolution of the bathymetry grid (Digital Elevation Model for seabed data following QUA-01-B GIS specifications) allows for enhanced visualization of depth and interesting seafloor features. The main use of the multibeam data is to reference the seismic profiles to the real seafloor (ZH Bathyelli v2 geoid).

In the acquired lines of the MED_AO6 Zone 4 offshore windfarm survey area, the depth range correspond between -86.78 meters in the shallowest part at the north-eastern region, to -112.85 meters in the deepest part at the southern region.

A colour table for the representation of the three-dimensional terrain model was created, from red -86.78 meters depth and magenta for the maximum depth -112.85 meters.



Figure 3-1: Colour table for the representation of the MBES terrain model.

Bathymetric data from the vessel multibeam sensor has been processed into a 1 meters grid size bathymetry for all the acquired lines.

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 23 of 49
Issue date	28/06/2023	Docun	nent uncon	trolled when printed/downloaded

	Project	Package	Issuer	Chrono	Revision	Status		
	MED_AO6	GPY	TEC	51	0	IFR		
	Title	MED_TEC_5	MED_TEC_51_Factual report - Seismic survey - OWF Zone 4 AO6 area_0					

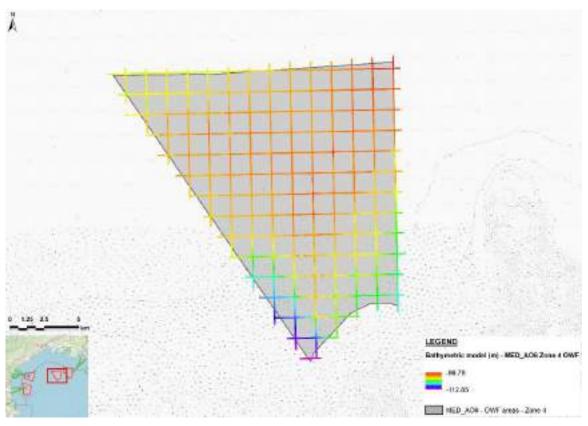


Figure 3-2: Whole bathymetric data grid model (1 x 1 m) for the MED_AO6 Zone 4 OWF.

Several examples of the results of the MBES processed data are shown below:

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 24 of 49
Issue date	28/06/2023	Docun	nent uncon	trolled when printed/downloaded

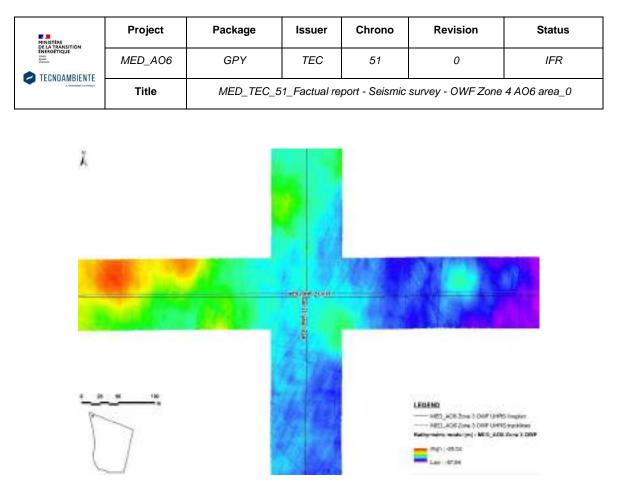


Figure 3-3: Detail of the bathymetric data grid model (1 x 1 m) for the offshore windfarm of the MED_AO6 Z3 area – Survey lines and tracklines MED_OWF_Z3_002-035.

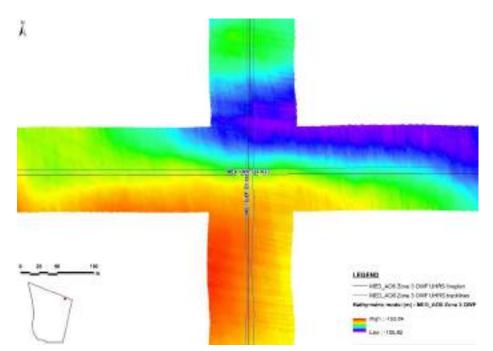


Figure 3-4: Detail of the bathymetric data grid model (1 x 1 m) for the offshore windfarm of the MED_AO6 Z3 area – Survey lines and tracklines MED_OWF_Z3_022-027.

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 25 of 49	
Issue date	28/06/2023	Document uncontrolled when printed/downloaded			

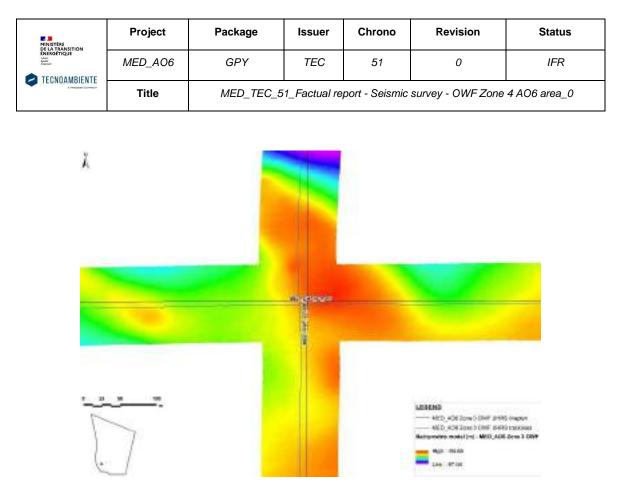


Figure 3-5: Detail of the bathymetric data grid model (1 x 1 m) for the offshore windfarm of the MED_AO6 Z3 area – Survey lines and tracklines MED_OWF_Z3_006-007.

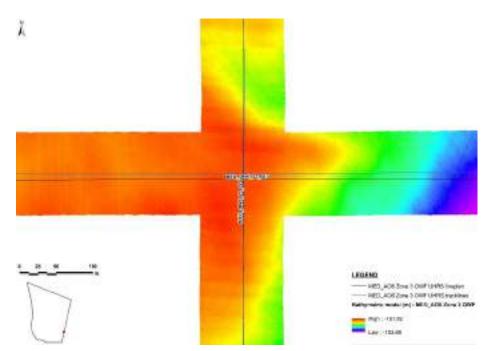


Figure 3-6: Detail of the bathymetric data grid model (1 x 1 m) for the offshore windfarm of the MED_AO6 Z3 area – Survey lines and tracklines MED_OWF_Z3_022-007.

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 26 of 49
Issue date	28/06/2023	Docun	nent uncon	trolled when printed/downloaded

	Project	Package	Issuer	Chrono	Revision	Status
	MED_AO6	GPY	TEC	51	0	IFR
	Title	MED_TEC_51_Factual report - Seismic survey - OWF Zone 4 AO6 area_0				

3.2. GEOLOGY

3.2.1. Geological setting from background data

The Gulf of Lions is located in the northwestern sector of the Mediterranean Sea bounded by the Pyrenees and the alps. It comprises a wide shelf and continental slope, before descending to the abyssal area of the Algero-Balearic Basin. The basin formed as a result of tectonic rifting during the Oligocene – Miocene period (Gorni, et al. 1994), leading to the accumulation of a large amount of clastic sediments forming a thick wedge on the inner shelf, and more than 2km on the outer shelf (Lofi, 2002). The continental shelf edge leads to the prograding margin observed in the Gulf of Lions during the end of the last glacial cycle. The geology within the Golf of Lion is described as a relatively low energy passive prograding margin, dominated by a rapid period of sedimentation during the Late Pleistocene, with layers of reworked sediments at a time when sea-levels were about 100m lower. At the end of the Last Glacial Maximum, sea levels were cyclically higher and lower as ice masses in the two hemispheres contracted and advanced. The deglacial succession overlies the major erosional discontinuity related sea level rises since the Last Glacial Maximum. It consists of basal transgressive deposits, subsequently reworked into dunes and sand ridges, interbedded with regressive prograding, marine derived sediments. The shelf 'relict' sands, pass rapidly into marine muds. The transition between sand and muds is outlined by a distinct regional step in sea-floor morphology.

3.2.2. Geological sequence

Within the depth of interest (up to 30m), the MED_AO6 OWF zone 4 area comprises a sedimentary succession of Late Pleistocene sediments, predominantly marine derived during periods of deposition. The shallow geology (30 m below seabed) within the MED_AO6 Zone 4 OWF site has been divided into units based on the sequence boundaries of major erosional surfaces. Five coherent stratigraphic packages, over the first 30 m subsea are distinguished, with the base of the units marked by sequence boundaries where there are clear unconformable, erosional events and a strong change in the dip angle of the

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 27 of 49	
Issue date	28/06/2023	Document uncontrolled when printed/downloaded			

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	MED_AO6	GPY	TEC	51	0	IFR
	Title	MED_TEC_5	1_Factual re	port - Seismic	survey - OWF Zone	4 AO6 area_0

sediments. The angle of dip may have been defined by the onlapping of sediments at the shelf edge, the change in a river pathway or other similar process. The presence of the strong change in dip suggests an erosional surface between that. These events mark clear sequence boundaries where there has been a change in conditions and hence separated into varying sedimentary units. The units are all interpreted to be similar in composition, and without geotechnical data to aid interpretation within zone 4 a succession of marine derived silty clay with some sand is interpreted.

Four horizons have been mapped within the Zone 4 AO6 OWF area which have been interpreted to represent, as already mentioned sequence variations. Some changes in the seismic acoustic character also suggests variations in the sediments within the units. The units are, the base of the uppermost marine drape (Base of Unit 1 - H05), a unit more chaotic in nature (base of Unit 2 - H15); the base of an acoustically quieter unit (Unit 3 - H20); and the base of a strongly dipping surface, with some evidence of channelling and coarser sediments within the channels.

Within the OWF, Unit 1, a blanket drape of acoustically quiet sediments, interpreted as a silty clay of Holocene age, and deposited since sea levels were at current levels which ranges in thickness from about 0.5 to 9 m (BSB). The base of unit 1 (H05) is interpreted as an erosional surface, with evidence across some areas of the OWF of channelling at the base, as well as a strong normal phase unconformity with parallel bedded reflectors above, interpreted as the drape of Unit 1, and a more chaotic acoustic character below (Unit 2). H05 has been mapped across the whole of the OWF, and converted into depths using an ASV of 1600m/s. An isopach to the depth to the base of Unit 1 is illustrated in Figure 3-9**¡Error! No se encuentra el origen de la referencia.** below as a thickness below seabed.

Below unit 1, a sand unit (Unit 2), is interpreted across much of the site. From the acoustic nature of the data, this unit appears to be more chaotic, and there is evidence of small channels, and much more variation in the amplitude response than found in Units 1 and 3. This suggests a variability in sediments found within the unit, with coarser lag deposits also

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 28 of 49
Issue date	28/06/2023	Docun	nent uncon	trolled when printed/downloaded

	Project	Package	Issuer	Chrono	Revision	Status		
	MED_AO6	GPY	TEC	51	0	IFR		
	Title	MED_TEC_5	MED_TEC_51_Factual report - Seismic survey - OWF Zone 4 AO6 area_0					

interpreted as present within the data. Unit 2 pinches out in the northwest of the area. In this area, the base of unit 1 H05 is interpreted to have eroded into the underlying Unit 3 sediments. Unit 2, where present, occurs between 0.8m below seabed and 21.5m below seabed, with the deepest section occurring in the centre of the survey area, below a sand bank. Unit 2 is illustrated in Figure 3-7, with the isopach to the depth to the base of Unit 2 illustrated on Figure 3-10.

Below Unit 2, there is a thin, acoustically quieter unit of sediments interpreted as marine derived silty clays (Unit 3). The base of this unit (H20) is gently dipping to the northwest. The base of this unit is present across most of the site, apart from where there are large channels incising into the base of it, as illustrated on Figure 3-8. The depth to the base of unit 3 varies between 6.1m below seabed in the north of the survey area, to 26m below seabed in the centre and southwest of the survey area. The acoustic character of the sediments infilling the channels suggest a different composition to the sediments found within Unit 3. An isopach to the base of Unit 3 is illustrated on Figure 3-11.

Unit 4 is interpreted as similar in nature to Unit 2. The base of Unit 4 (H30) is marked by a strong variation in the angle of dip and the acoustic character of the sediments. Within Unit 4, there are predominantly parallel bedded reflectors with these in the north becoming acoustically more chaotic and less coherent. This suggests a degree of weathering or erosion within this unit, interpreted to have occurred during a period of reduced sea levels when the sediments were aerially exposed. The depth to the base of Unit 4 is illustrated on Figure 3-12 as an isopach, and occurs between 11.4m below seabed in the northwest of the OWF, and at 37m below seabed in the southeast, or is absent where the large channels are observed in the seismic data in the south

Below Unit 4, Unit 5 is interpreted to be similar in acoustic character to Unit 4, the primary difference between the two units is the angle of the dip, and hence the sequence boundary between the two units representing a change in environmental conditions during their deposition.

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 29 of 49	
Issue date	28/06/2023	Document uncontrolled when printed/downloaded			

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	MED_AO6	GPY	TEC	51	0	IFR		
	Title	MED_TEC_5	MED_TEC_51_Factual report - Seismic survey - OWF Zone 4 A06 area_0					

Channels are interpreted in the south of the survey area. The location of these are shown on Figure 3-8, and Figure 3-13 displays their extents. These are predominantly beyond the depth of interest for anchor emplacement but have been included for completeness due to their size.

Confidentiality	Diffusion restrein	te (restricted)	Pages	Page 30 of 49
Issue date	28/06/2023	Docun	nent uncon	trolled when printed/downloaded

MINISTÈRE DE LA TRANSITION ENERGÉTIQUE	Project	Package	Issuer	Chrono	Revision	Status		
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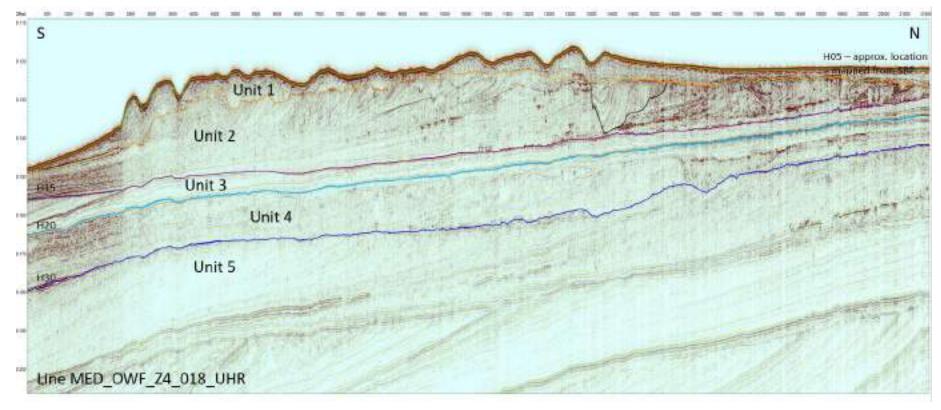


Figure 3-7: Line AO6_OWF_Z4_018.

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 31 of 49
Issue date	28/06/2023	Document uncontrolled when printed/download		nent uncontrolled when printed/downloaded

MINISTÈRE DE LA TRANSITION ENERGÉTIQUE	Project	Package	Issuer	Chrono	Revision	Status		
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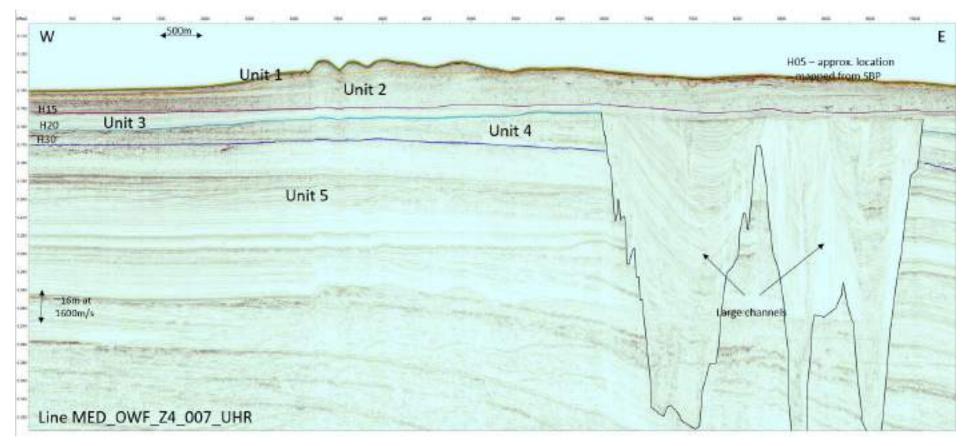


Figure 3-8: Line AO6_OWF_Z4_007.

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 32 of 49
Issue date	28/06/2023		Docun	nent uncontrolled when printed/downloaded

MINISTÈRE DE LA TRANSITION	Project	Package	Issuer	Chrono	Revision	Status	
	MED_AO6	GPY	TEC	51	0	IFR	
	Title	MED_TEC_51_Factual report - Seismic survey - OWF Zone 4 AO6 area_0					

Table 4: Shallow geological units.

Unit	Upper surface	Lower surface	Description	Depositional Environment
1	Seabed	H05	Uppermost unit mapped off the Innomar data, with an erosional surface at the base marked by small channels. Thins to 0.9m in the northwest and is up to 9m thick in the center of the survey area. The erosional base suggests there may be some coarser or lag deposits present	Shallow marine, a drape of sediment deposited since sea level rise and the area was exposed.
2	H05	H15	Discontinuous reflectors, a package of sediment marking multiple events of depositional reworking and erosion. Exposed above sea level. Acoustically of higher amplitude. Evidence of unconformities, channelling within. Strong normal phase unconformable reflector at the top and base of the unit, interpreted as predominantly sandy in nature, some coarser material, and occasional small clay beds may be present within.	Estuarine/lacustrine depositional and some terrestrially reworked
3	H15	H20	Acoustically quiet unit with little to no structure within it. Interpreted as marine clays rapidly deposited during a period of higher sea levels.	Marine deposited parallel bedded acoustically quieter unit.
4	H20	H30	Discontinuous reflectors, a package of sediment marking multiple events of depositional reworking and erosion. Exposed above sea level. Acoustically of higher amplitude. Evidence of unconformities, channelling within. Strong normal phase reflector at the top and base of the unit, interpreted as predominantly sandy in nature, some coarser material, and occasional small clay beds may be present within.	Estuarine/lacustrine depositional and some terrestrially reworked
5	H30		Discontinuous reflectors, a package of sediment marking multiple events of depositional reworking and erosion. Exposed above sea level. Acoustically of higher amplitude. Evidence of unconformities, channelling within. Strong normal phase unconformable reflector at the top of the unit. Interpreted as predominantly sandy in nature, some coarser material, and occasional small clay beds may be present within.	Marine deposited parallel bedded acoustically quieter unit.

Confidentiality	Diffusion restrein	e (restricted)	Pages	Page 33 of 49
Issue date	28/06/2023	Docum	nent uncon	trolled when printed/downloaded

MINISTÈRE DE LA TRANSITION	Project	Package	Issuer	Chrono	Revision	Status	
	MED_AO6	GPY	TEC	51	0	IFR	
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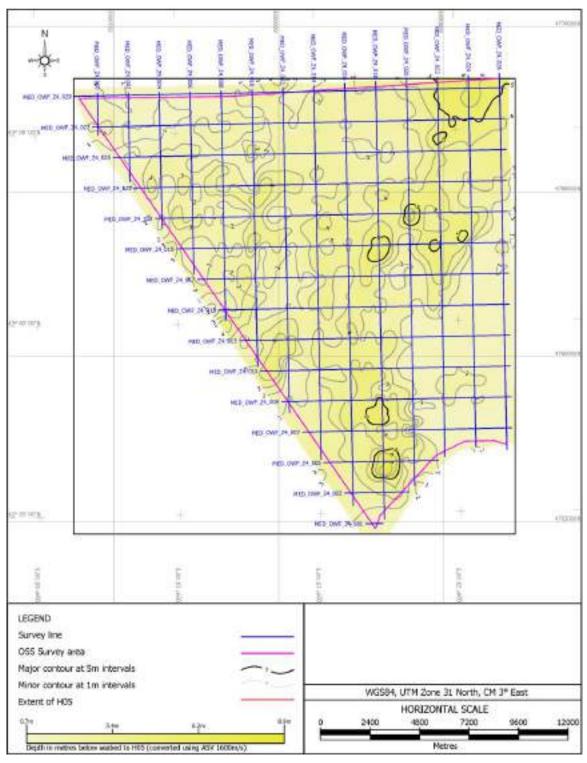


Figure 3-9: Isopach of depth to Base Unit 1 (H05)

Confidentiality	Diffusion restreint	e (restricted)	Pages	Page 34 of 49
Issue date	28/06/2023	Docun	nent uncon	trolled when printed/downloaded

MINISTÈRE	Project	Package	Issuer	Chrono	Revision	Status
	MED_AO6	GPY	TEC	51	0	IFR
	Title	MED_TEC_5	1_Factual re	port - Seismic	survey - OWF Zone	4 AO6 area_0

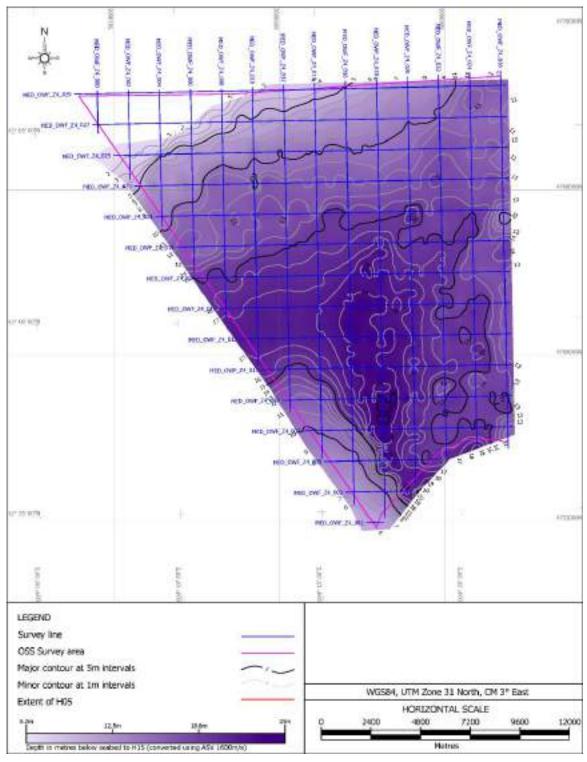


Figure 3-10: Isopach of depth to Base Unit 2 (H10).

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 35 of 49
Issue date	28/06/2023	Document uncontrolled v		trolled when printed/downloaded

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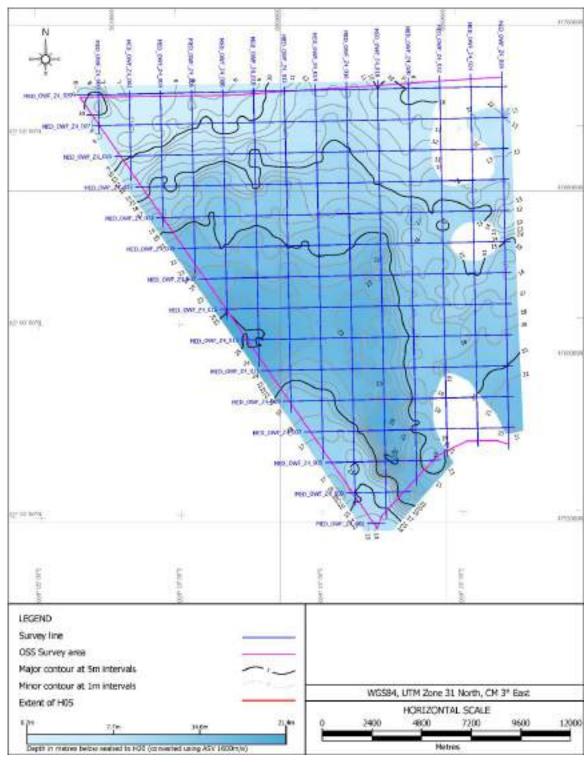


Figure 3-11: Isopach of depth to Base Unit 3 (H20).

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 36 of 49
Issue date	28/06/2023	Document uncontrolle		trolled when printed/downloaded

MINISTÈRE	Project	Package	Issuer	Chrono	Revision	Status
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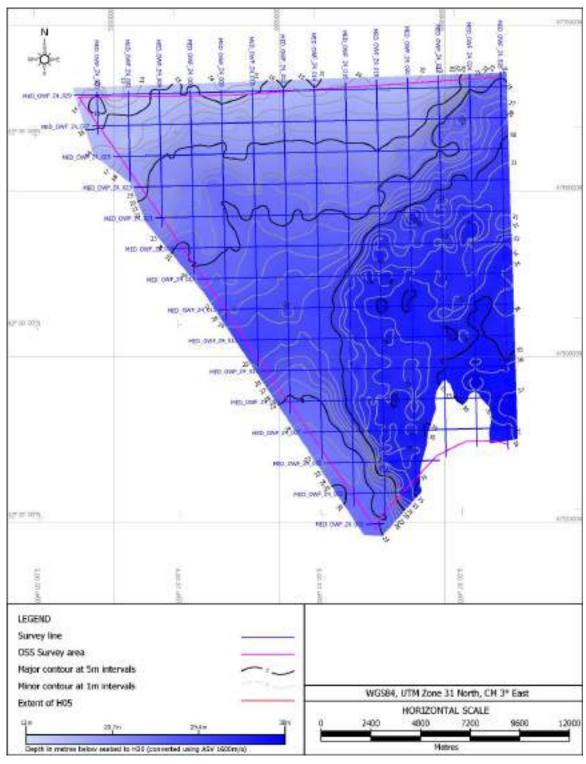


Figure 3-12: Isopach of depth to Base Unit 4 (H30).

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 37 of 49
Issue date	28/06/2023	Document unco		trolled when printed/downloaded

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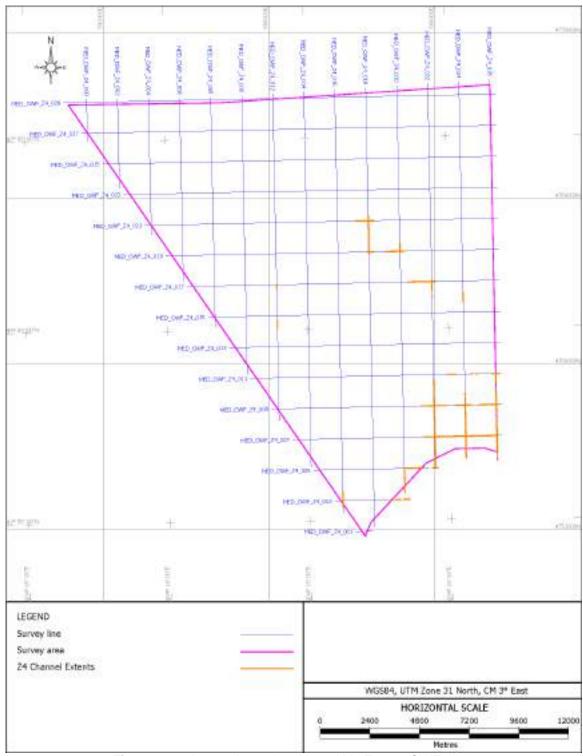


Figure 3-13: Large channel extents within Zone 4 OWF area.

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 38 of 49
Issue date	28/06/2023	Docun	nent uncon	trolled when printed/downloaded

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3.2.3. Geohazards

The shallow geology has been checked for any evidence of any shallow geohazards that may affect the installation or operation of a floating wind farm. Constraints may relate to composition and distribution variability of sediments (at the seabed and in the subsurface) within the first 30 m below the seafloor. Other constraints may relate to past or presently active geological processes, such as gas seepage and faulting.

A summary of geological conditions and potential constraints on infrastructure and engineering activities, applied to the Zone 4 MED_AO6 site, is provided in (Table 5) modified after (Mellet, Long, Carter, Chiverell, & Van Landeghem, 2015).

Within the upper 30m of the Zone 4 OWF, there are various conditions, that may or may not be considered a hazard based on the installation design.

- Large channels observed in the south of the survey area within the upper 30m will mean there are sediment variations which would need investigating before anchor emplacement is planned.
- Coarse or lag deposits interpreted to be present in the base of Unit 1 and throughout Unit 2 may need to be considered for anchoring purposes.

There is no evidence of gas, or slope failure or other similar geohazards within the Zone 4 OWF, with a full list of all geohazards and the potential constraint listed in Table 5 below.

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 39 of 49
Issue date	28/06/2023	Docun	nent uncon	trolled when printed/downloaded

MINISTÈRE DE LA TRANSITION	Project	Package	Issuer	Chrono	Revision	Status			
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Table 5: Geological characteristics / processes and potential constraints.

Geological characteristic / process	Geological Potential constraint	
	Seabed sediments	
Soft muds	Low strength means they will not bear large loads	Probable
Coarse lag (gravel to boulders) deposits		
	Mobile sediment	
Migrating bedforms change topography (can create seabed features several metres height)	Can bury or expose structures or create a barrier to activities	Not evident
Mobile sediment can change sediment characteristics at seabed	Mobile sediment is constantly changing. Therefore, expect variation between samples taken from the same site at different times.	Not evident
Bedforms can migrate in the opposite direction to that predicted from morphology and tidal residual	Do not assume sediment migration pathways from morphology. Repeat bathymetric surveys should be carried out.	Not evident
	Gas/fluid escape and MDAC	
Gas or fluid present in shallow subsurface	Can lead to blowouts when drilling	Not evident
MDAC	Creates a hard substrate that is recognized as a special habitat that must be preserved	Not evident
	Quaternary	
Variable sediment thickness	Locally, sediment thickness can change from thin (<5 m) to thick (> 50 m) over a short distance	Expected
Variable lithology (vertically and spatially)	Glacial processes rework and deposit sediments that are highly variable over large areas. Smaller and isolated features such as channels are not always mapped across the site. With the large line spacing some of the variations may not be incorporated into the interpretation.	Expected
Heterogeneous sediment composition	Sediments are typically glacially diamict which are poorly sorted mixtures of silt, sand, gravel, and clay. Diamicts can be interbedded with sands.	Probable
	Bedrock	
Bedrock outcrop at seabed	Provides a hard substrate for emplacement of seabed infrastructure.	Not evident
Faulting	Active faults are susceptible to ground surface ruptures that can compromise infrastructure; seabed forms that indicate pre-existing seabed instability, surface displacements, or fluid escapes are conditions that pose risk to infrastructure; Sub- surface fault zones may provide preferential conduits for gas migration, or may be hydraulically active during (or shortly after) earthquakes	Not evident

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 40 of 49
Issue date	28/06/2023	Document uncontrolled when printed/downloaded		trolled when printed/downloaded

MINISTÈRE DE LA TRANSITION	Project	Package	Issuer	Chrono	Revision	Status		
	MED_AO6	GPY	TEC	51	0	IFR		
	Title	MED_TEC_51_Factual report - Seismic survey - OWF Zone 4 AO6 area_0						

The depth values were converted from time (TWT) using sound velocity of 1,600 m/s in sediments and 1,510 m/s in the water column.

3.2.4. Background data summary and regional geology

To provide background context for regional stratigraphy and structural geology relevant portions of text from (Bassetti, 2006) are reproduced here:

In the Gulf of Lions, which is considered as a relatively low energy continental shelf, most of the authors still consider that the offshore sands are relict features, only the transgressive processes, at a time when sea-level was lower by about 100 m, being able to rework sediments (Aloïsi, 1986; Berné et al., 1998; Monaco, 1971). However, ultra-high resolution seismic data, coring and 14C dating, as well as numerical modelling of wind stress on oceanic circulation, allow us to demonstrate that a mobile carpet of sand is periodically active at the shelf edge, feeding slope and rise deposits and contributing to the episodic reworking of shelf morphology.

Morphology and seismic facies of post-glacial deposits

(a) Sand ridges- In the studied area, the major morphological feature is represented by the sand ridges, localized between 95 and 110 m water depth. They have limited areal (Bassetti et al. (Marine Geology)) distribution, variable heights (up to 9 m) a mainly WNW-ESE orientation, as recognized on the bathymetric maps. They have an irregular topography and mainly show a linear, elongated shape. On the chirp profiles their surface is smooth, they form bodies of maximum length of 5 km and they rest on a major erosional surface (ravinement surface) that is possibly exposed beyond the ridge field. The ridges have an asymmetric transverse profile (with the steepest slope facing the SW). At times, they show a nearly symmetric profile, but it concerns only the smaller bodies. They show distinct clinoforms, dipping in the SW direction and some chaotic internal reflections probably in relation with coarse-grained material diffracting seismic waves. However, some major erosional surfaces (discontinuities) can be recognized within the ridge, that may be related

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 41 of 49
Issue date	28/06/2023	Document uncontrolled when printe		trolled when printed/downloaded

MINISTÈRE DE LA TRANSITION	Project	Package	Issuer	Chrono	Revision	Status		
	MED_AO6	GPY	TEC	51	0	IFR		
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to storm events affecting the ridge growth, although they cannot be correlated from one ridge to another.

(b) the dunes were only detected when we used high-resolution swath bathymetric systems, such as the EM 1000 and EM1002S. They have an average spacing of 130 m and maximum height of 2 m. Their great axis has a NNW-SSE orientation in the NE part of the surveyed area, turning progressively to NW-SE in the SW corner. Their internal structure was not detectable considering their small size. They are classified as transverse dunes in the sense of Ashley (1990) and they clearly rework the shape of the sand ridges. On top of the sand ridges, chirp and sub-bottom seismic profiles display a thick pattern of parallel reflections, that was first considered as the result of some ringing effect representing the pulse length of the seismic sources, instead of a real sedimentary layer. However, extensive coring and bathymetric data demonstrated that a distinct layer actually exists at the sea-floor interface.

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 42 of 49
Issue date	28/06/2023	Document uncontrolled when printed/downloaded		trolled when printed/downloaded

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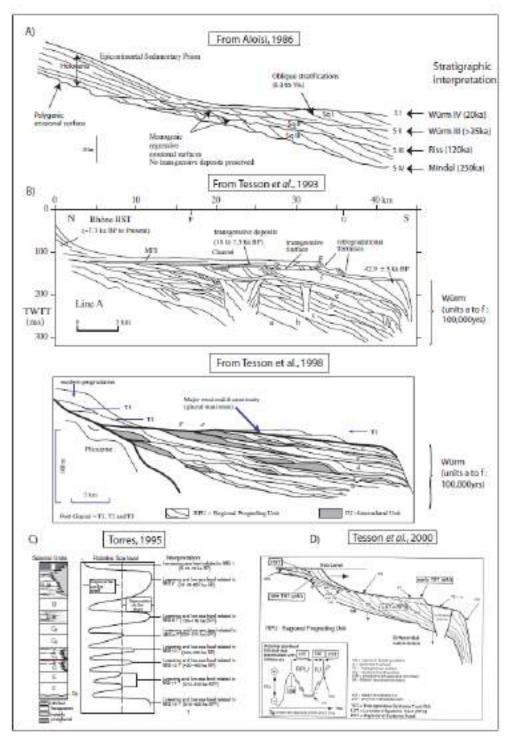


Figure 3-14: Figure showing the geology (Figure 3 from Rabineau et al., 2003; Benabdellouahed, 2011; Paquet et al. - b, in preparation).

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 43 of 49
Issue date	28/06/2023	Docum	nent uncon	trolled when printed/downloaded

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3.2.5. Conclusions and recommendations/comments

Horizons were mapped to define units of similar sedimentary facies based on acoustic nature and known geology and environmental conditions during the time of deposition from background material. These have been illustrated as isopachs to show rough sedimentary thicknesses and assist with a ground model. The sediment types and any hazards present were mapped.

No evidence of shallow gas or faulting is observed.

To improve knowledge of potential geological constraints the following data acquisition could prove beneficial.

- 1. Acquire additional high-resolution seismic data at a closer line spacing to improve spatial mapping of stratigraphic units.
- 2. Acquire sidescan sonar imagery data to coincide with the multibeam bathymetry and backscatter data.
 - a. A detailed seafloor mapping with sidescan sonar data will identify potential natural and anthropogenic seafloor geohazards.
- 3. Acquire repeat multibeam bathymetry and backscatter data during sidescan sonar data acquisition
 - a. Use this comparative multibeam bathymetry data to assess potential for sediment mobility.
 - Although a recent study concluded that sediment mobility is not apparent at the NOR_AO4 OSS site, monitoring for its potential could be beneficial for long-term development planning.
- 4. Acquire seabed ground-truthing "light" geotechnical data (e.g., grab samples) to confirm the variable seafloor composition illuminated with the sidescan sonar data.

Confidentiality	Diffusion restreinte (restricted)		Pages	Page 44 of 49
Issue date	28/06/2023	Document uncontrolled when printed/downloaded		trolled when printed/downloaded

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4. REFERENCES

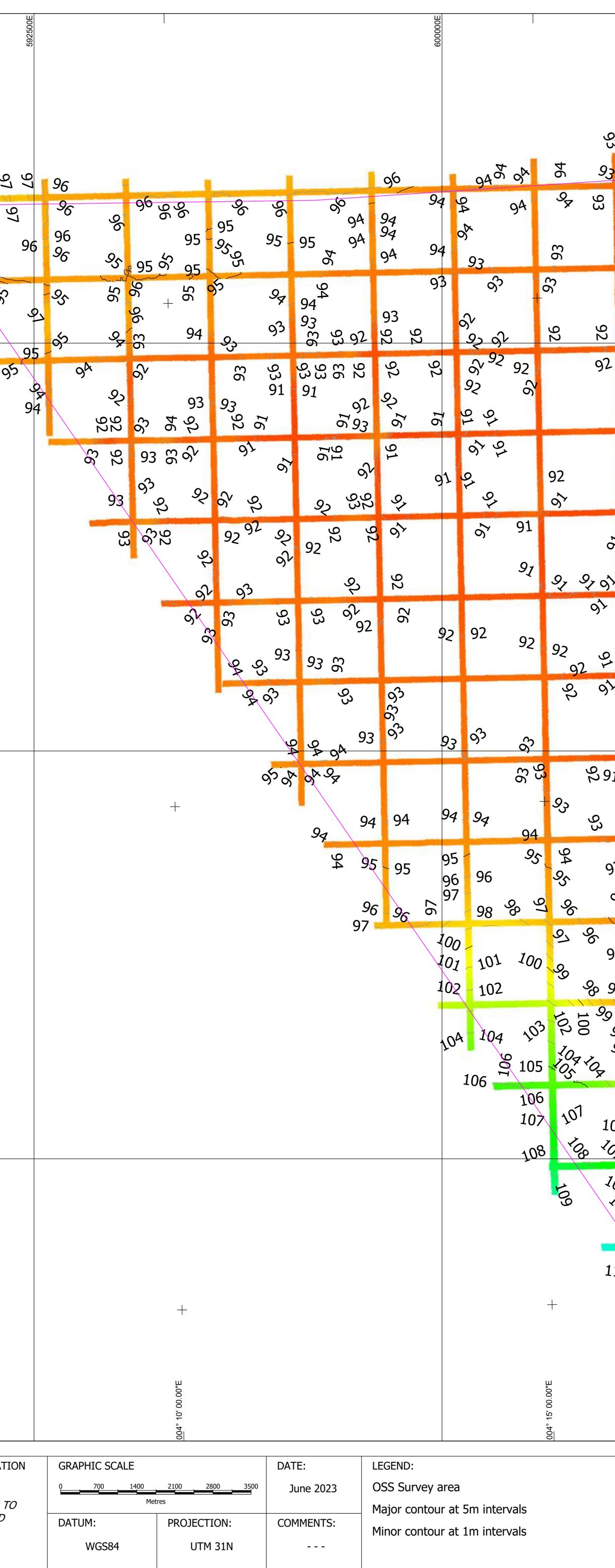
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Confidentiality	Diffusion restreinte (restricted)		Pages	Page 45 of 49
Issue date	28/06/2023	Docun	nent uncon	trolled when printed/downloaded

APPENDIX I – CHARTING

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2	Z4_OWF_GRADIENT
3	Z4_OWF_PROFILE_018
4	Z4_OWF_PROFILE_023
5	Z4_OWF_PROFILE_023
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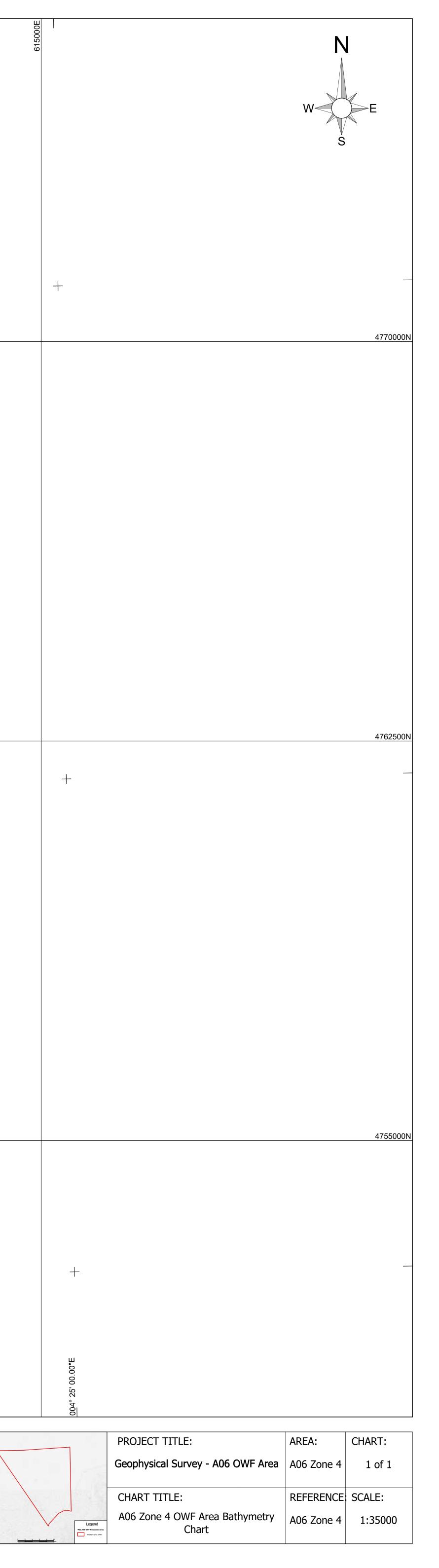
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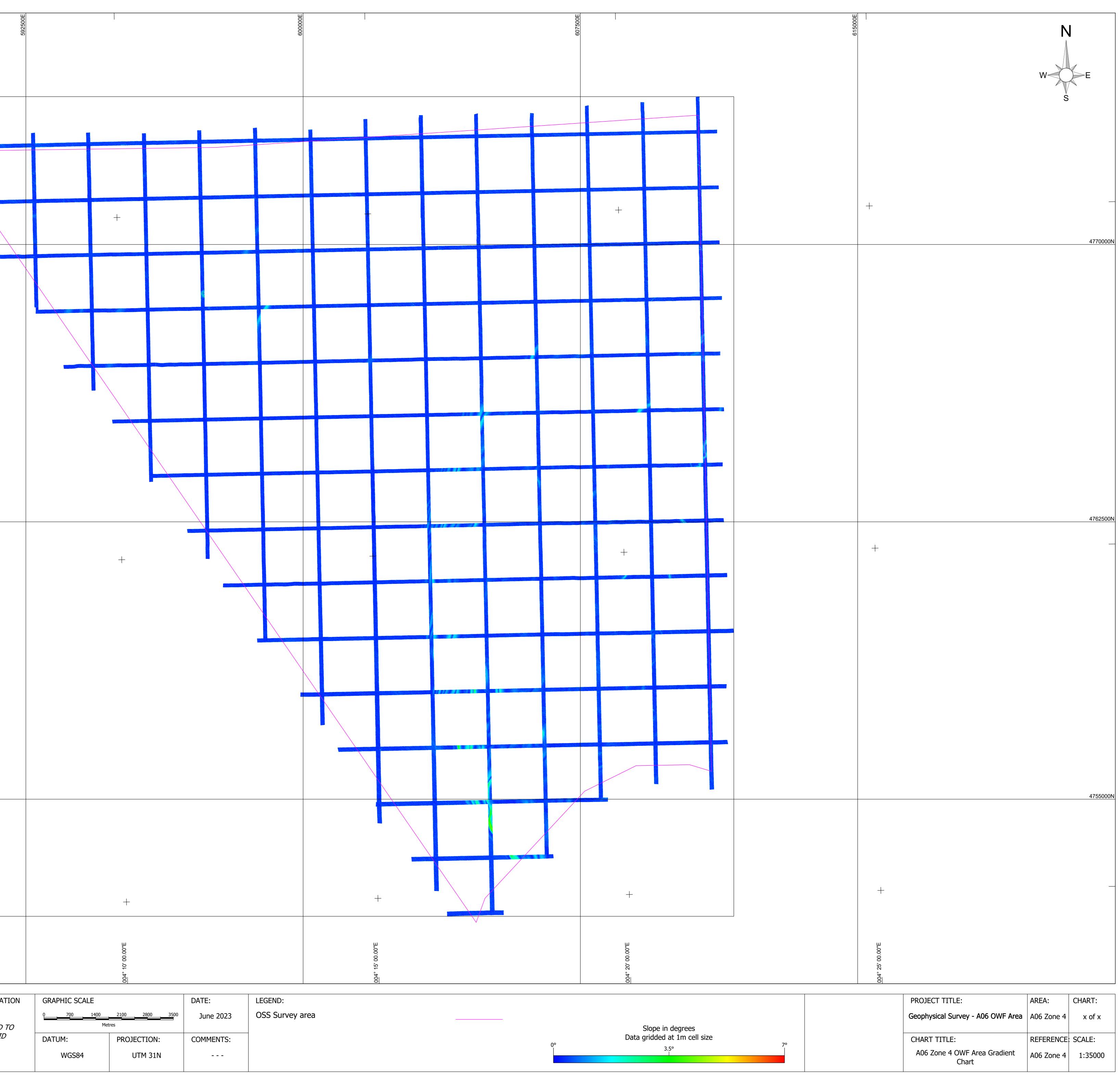
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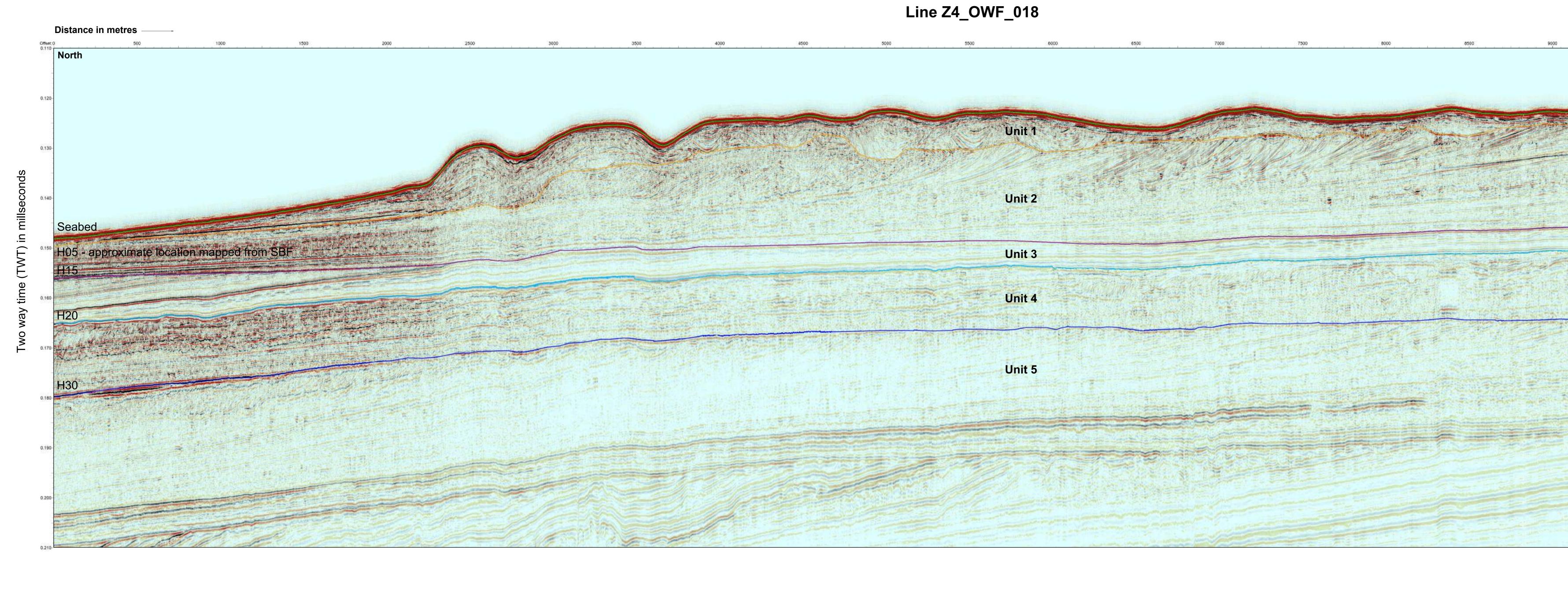
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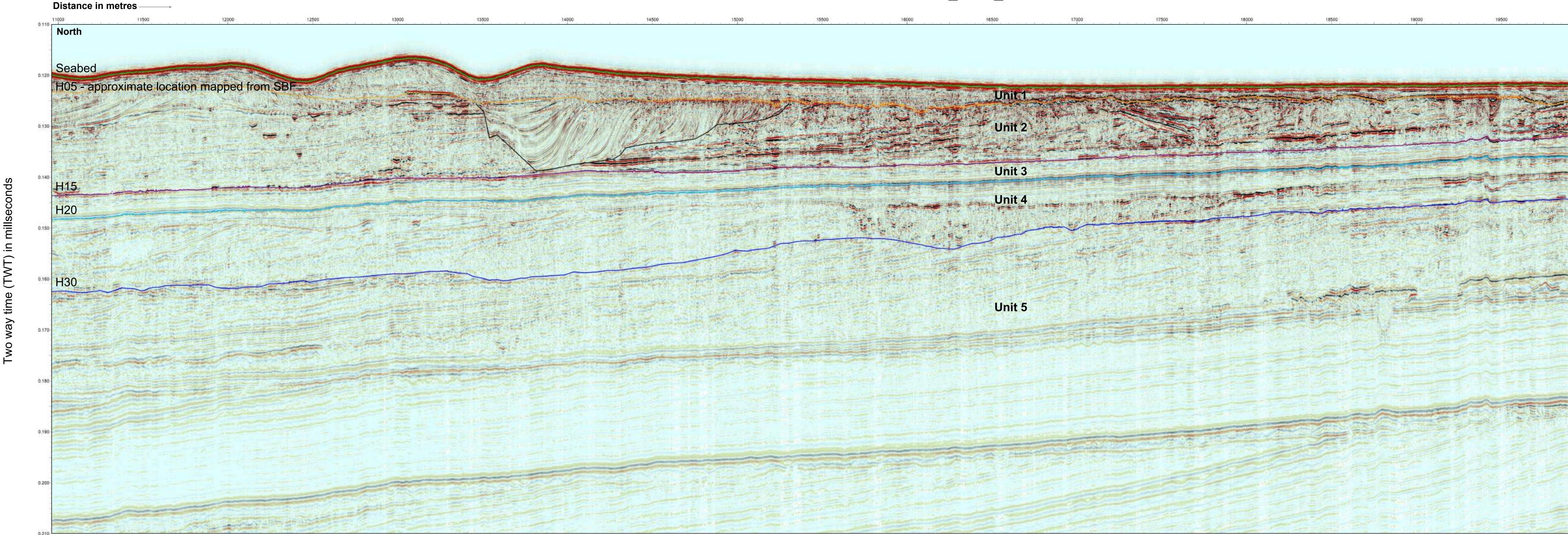
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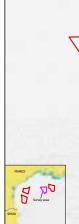


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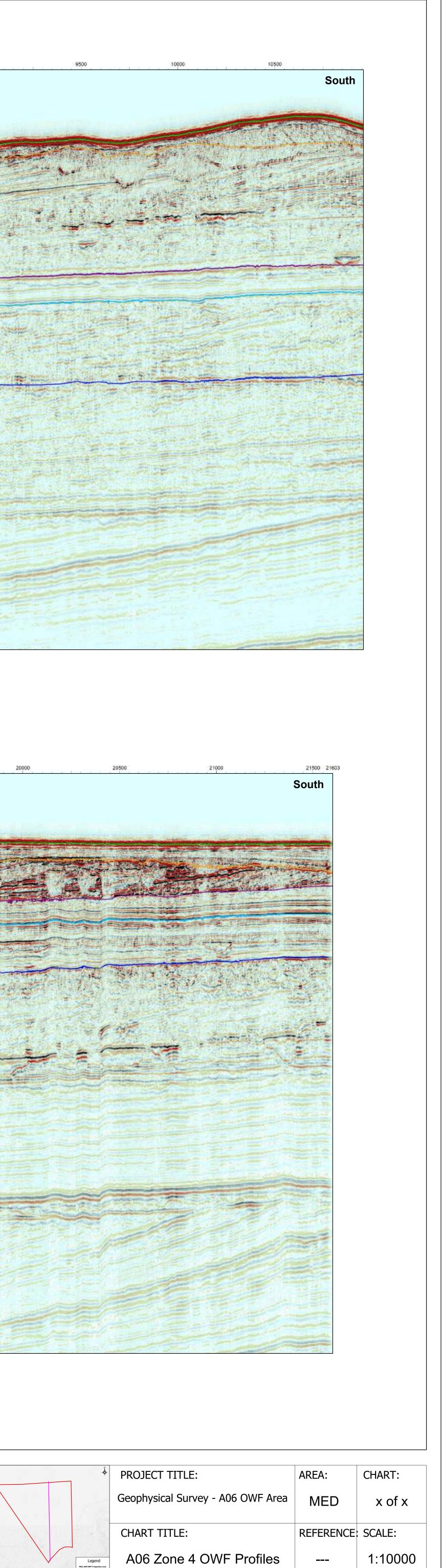
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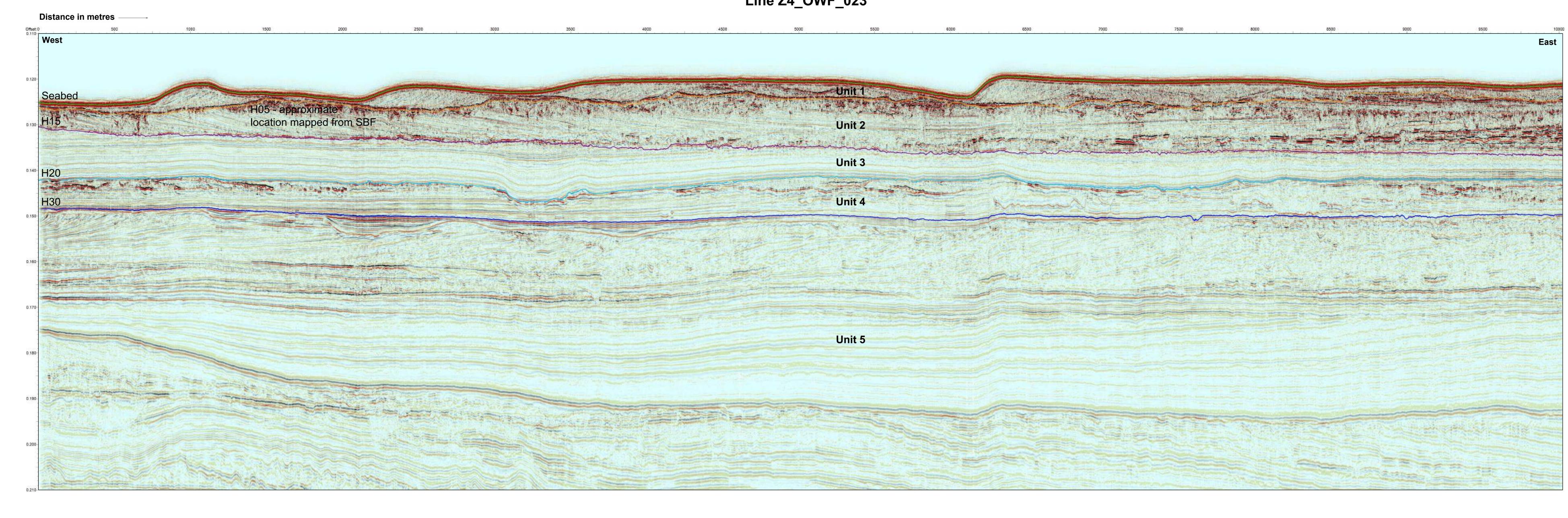
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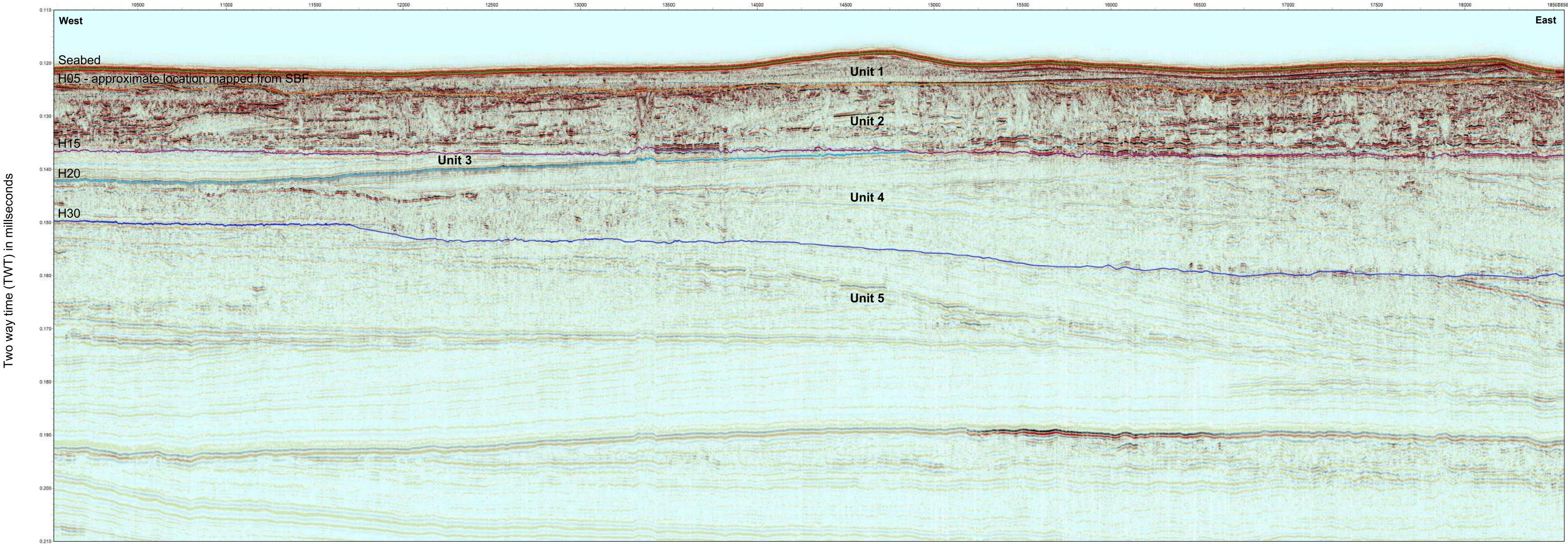
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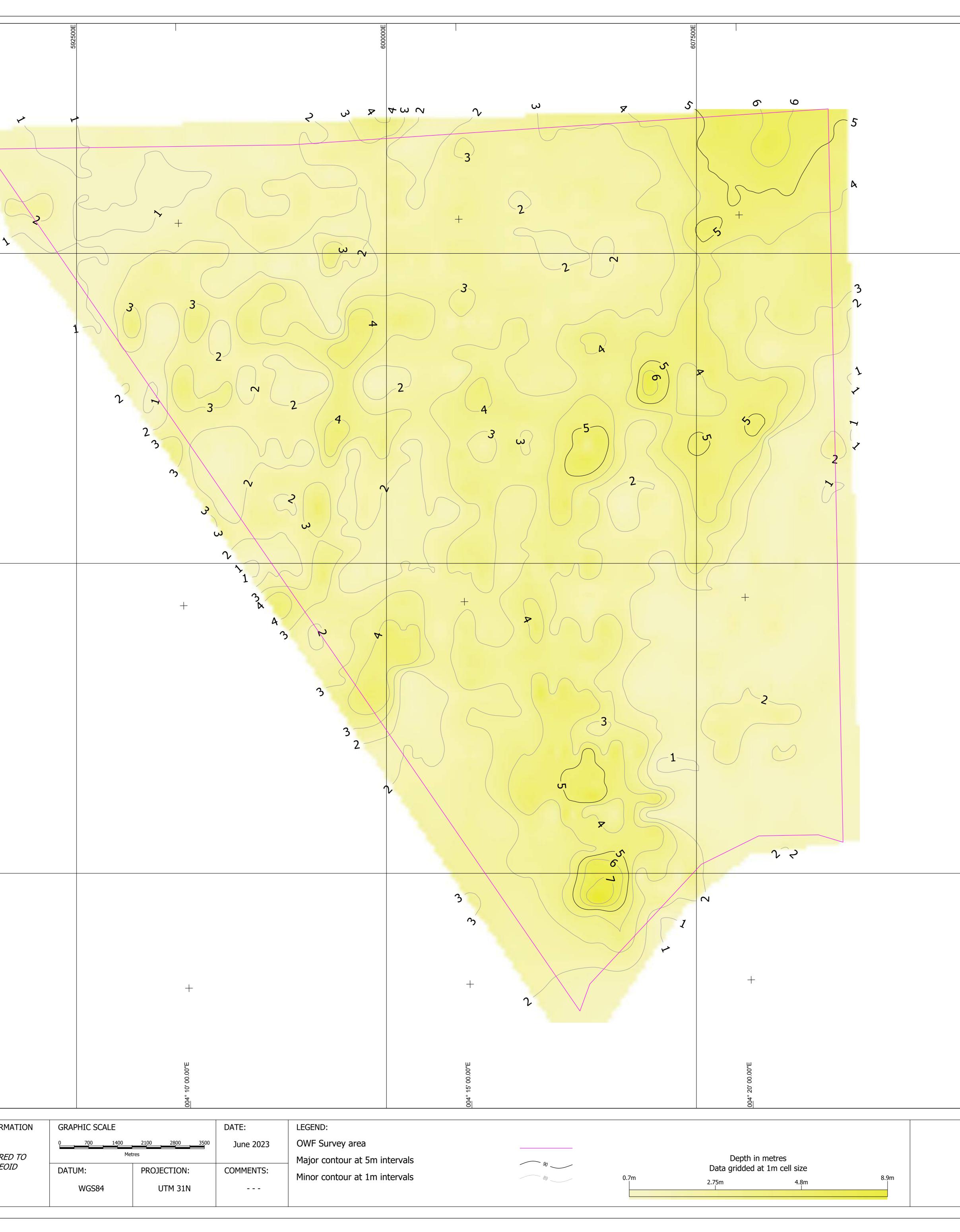


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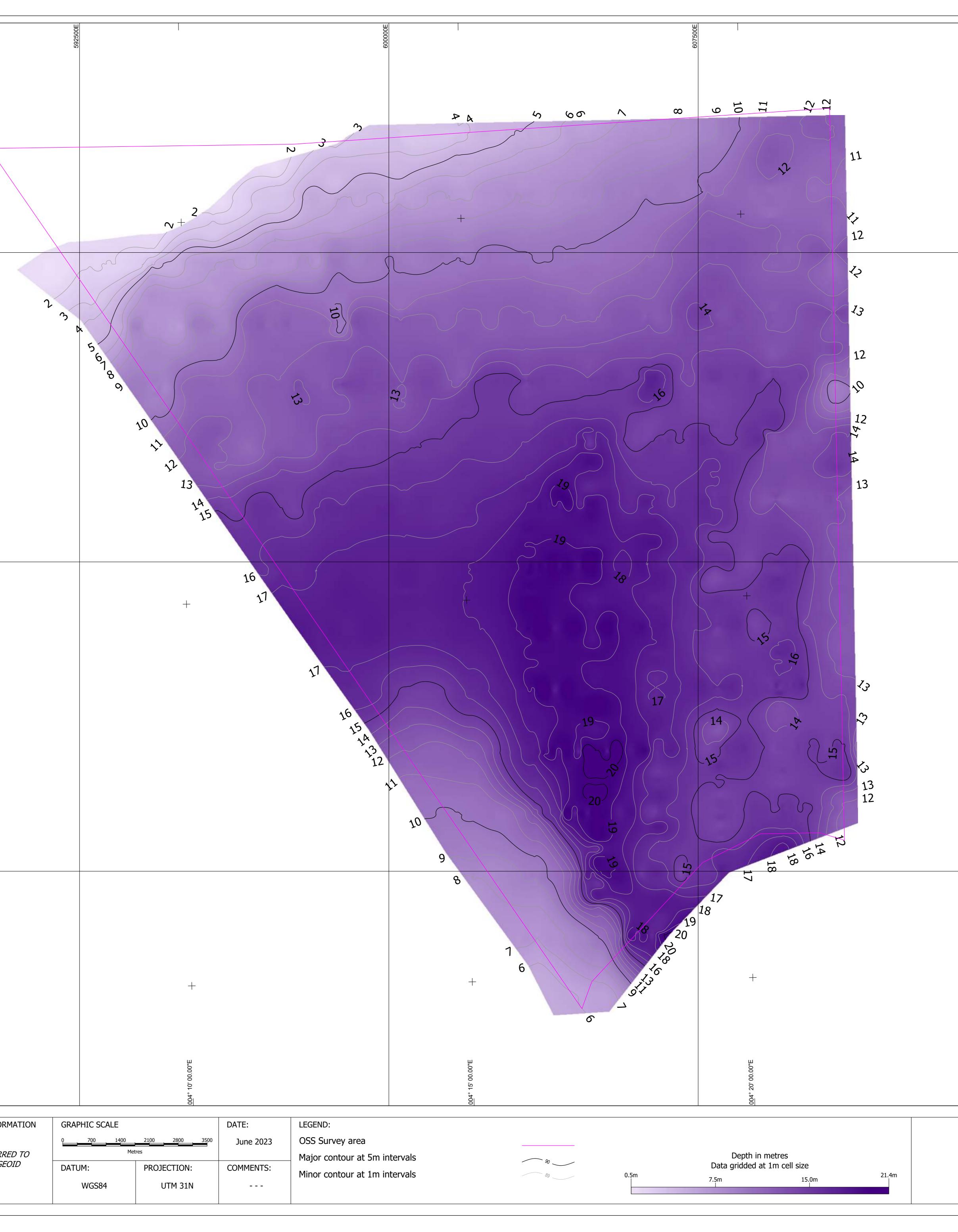
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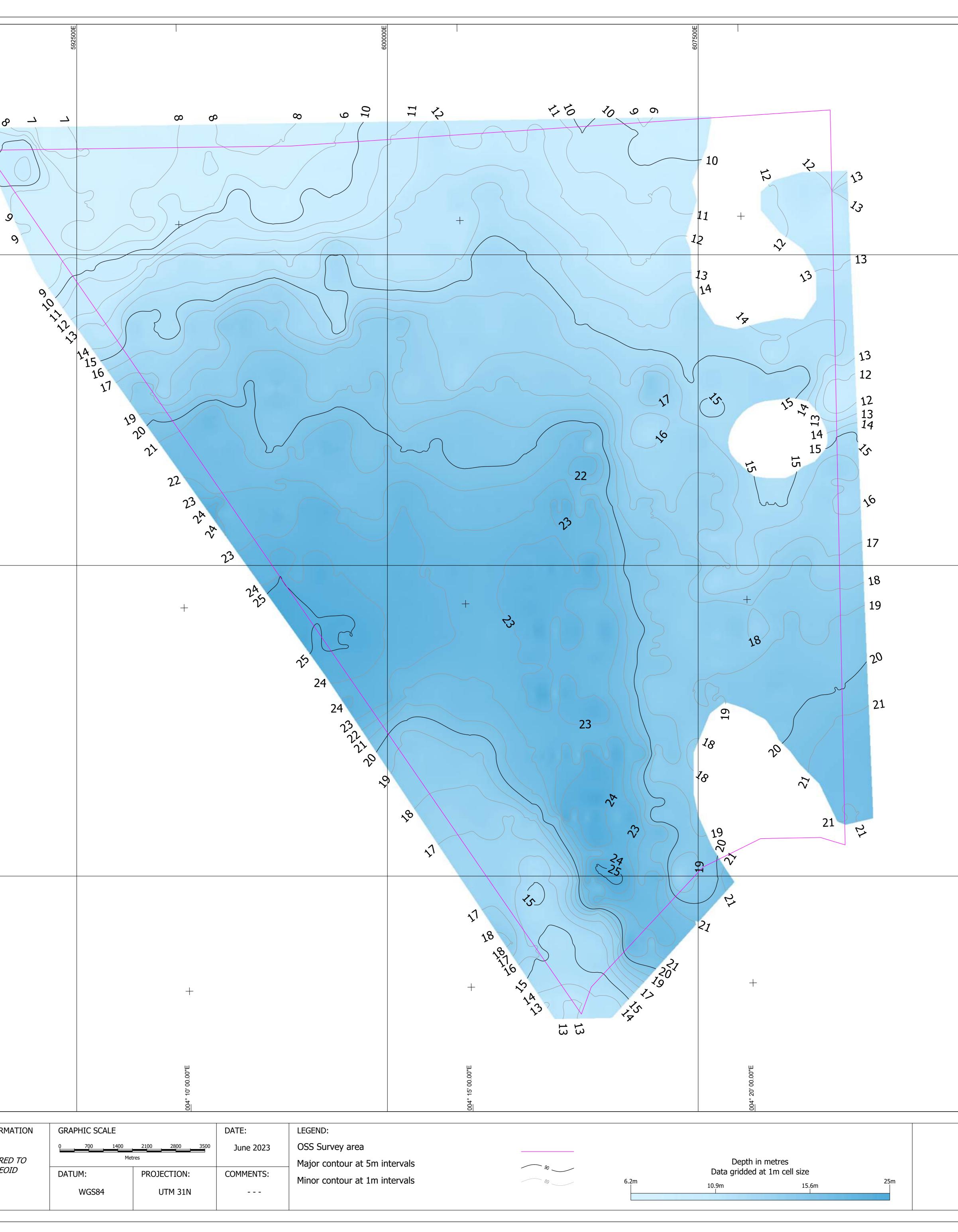
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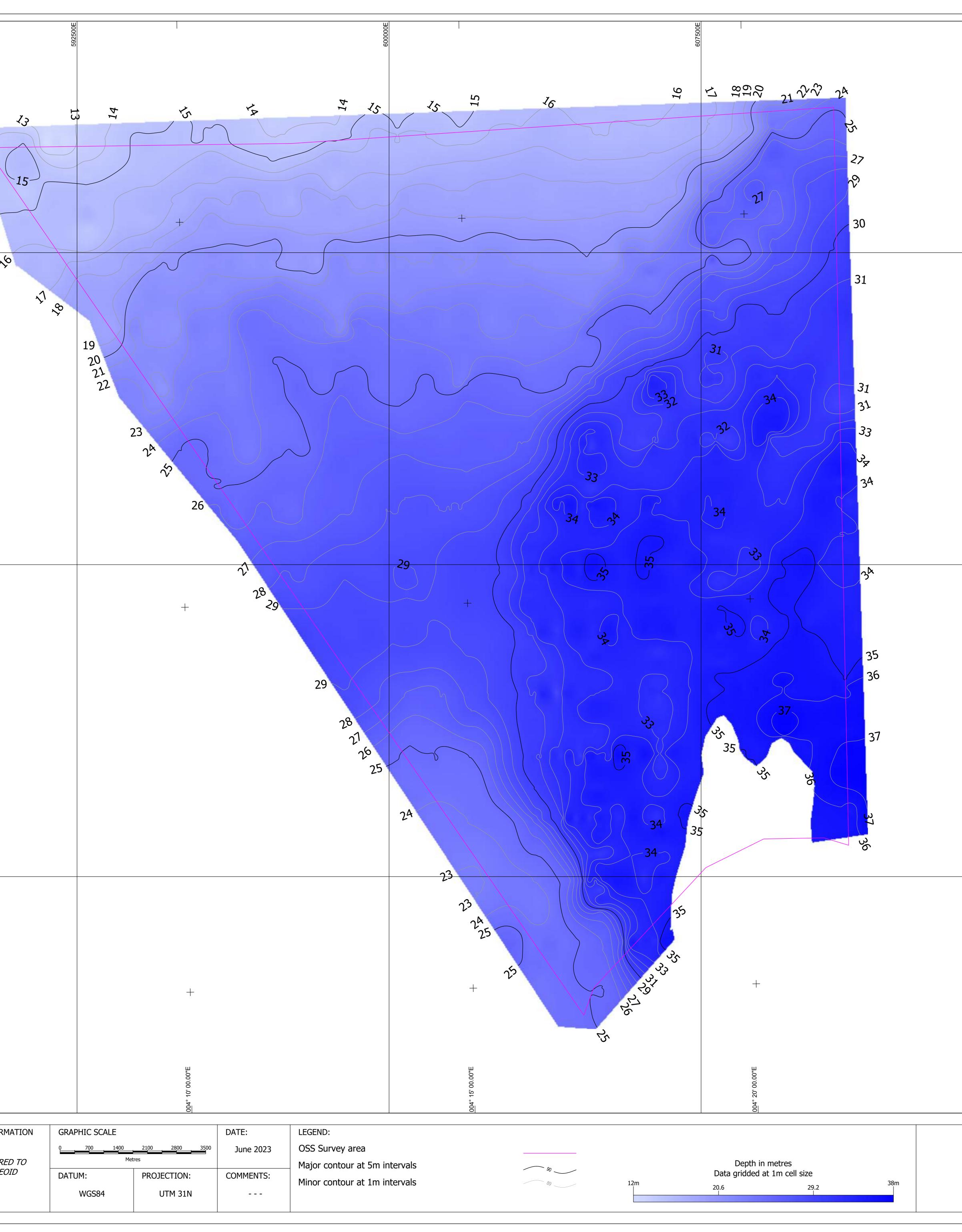
615000E		N W S	E
	_		4770000N
	+		<u>4762500N</u>
			4755000N
	+		
	004° 25' 00.00"E		
	PROJECT TITLE:	AREA: A06 Zone 4 REFERENCE	CHART: 1x of 1 SCALE:

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<u>43</u> ° 00' 00.00"N <u></u>			+	
<u>42</u> ° 55' 00.00"N			+	
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				CARTOGRAPHIC INFORM
MINISTÈRE DE LA TRANSITION ÉCOLOGIQUE		ECNOAMBIENTE		ELEVATION REFERRE BATHYELLI v2 GEC



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MINISTÈRE DE LA TRANSITION ÉCOLOGIQUE	6	TECNOAMBIENTE		ELEVATION REFERRE BATHYELLI v2 GEO
ÉCOLOGIQUE				



615000E	N w t E s		
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	+		<u>4762500N</u>
			<u>4755000N</u>
	+		
	PROJECT TITLE: Geophysical Survey - A06 OWF Area	AREA: A06 Zone 4	CHART: 1x of 1
	CHART TITLE: A06 Zone 4 OWF Area Isopach H30 Chart	REFERENCE A06 Zone 4	

APPENDIX II – SEISMIC PROCESSING OVERVIEW



Le réseau de transport d'électricité



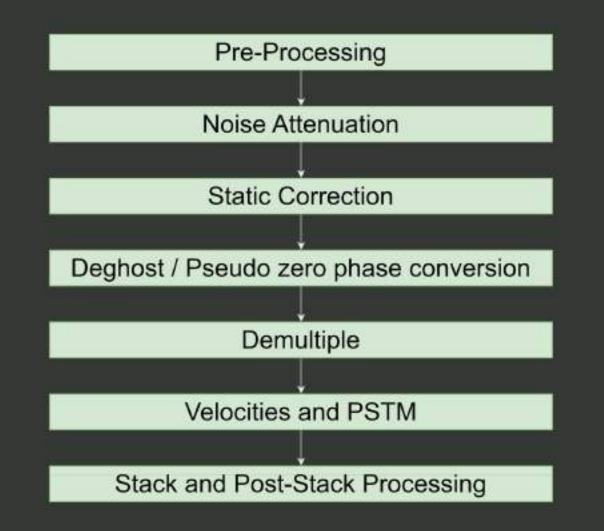
MED_AO6 SEISMIC PROCESSING OVERVIEW

UHR SURVEY PARAMETERS

PARAMETER	VALUE	PARAMETER	VALUE
Sample Rate	0.0625ms	Active Streamer Length	75m
Record Length	0.250ms	Number of channels	48
Shot Point Interval	Im	Group Length	Channels 1-24 :1m Channels 25-48 : 2m
Source	Sparker – GSO – 400 tips	Target Tow Depth	lm +/-0.5m
Target Source Tow Depth		Near Offset	~5-6m

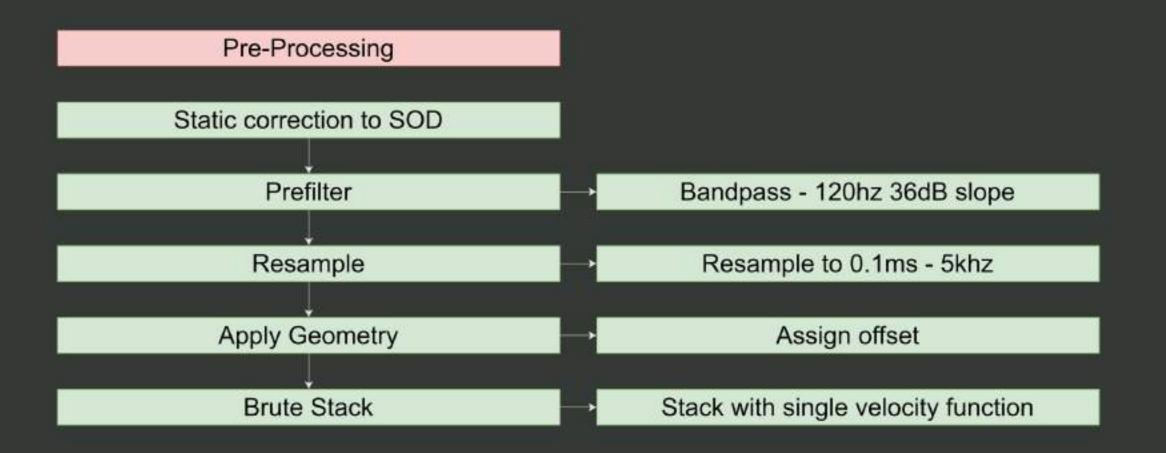


SEISMIC PROCESSING OVERVIEW



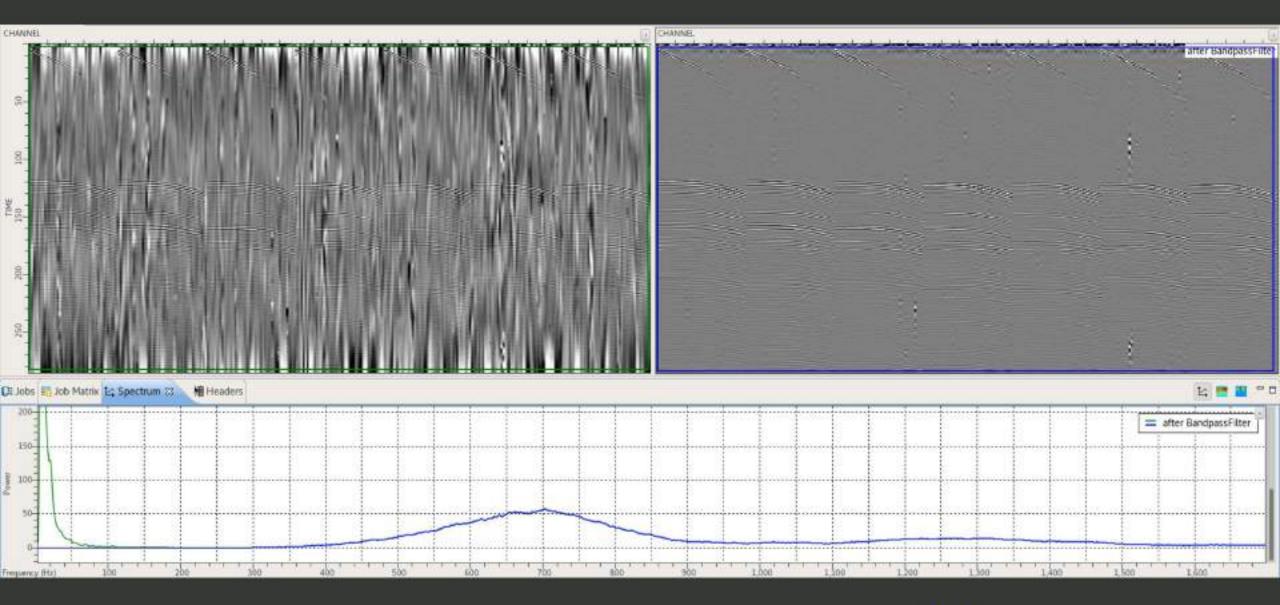


PREPROCESSING



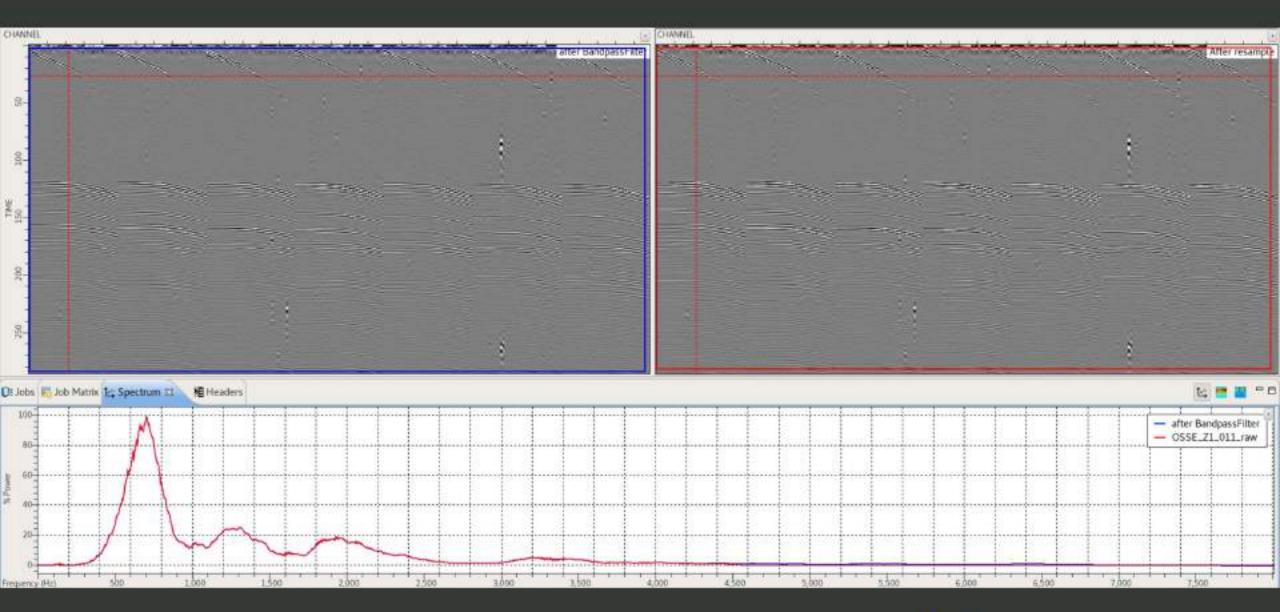


PRE-FILTER



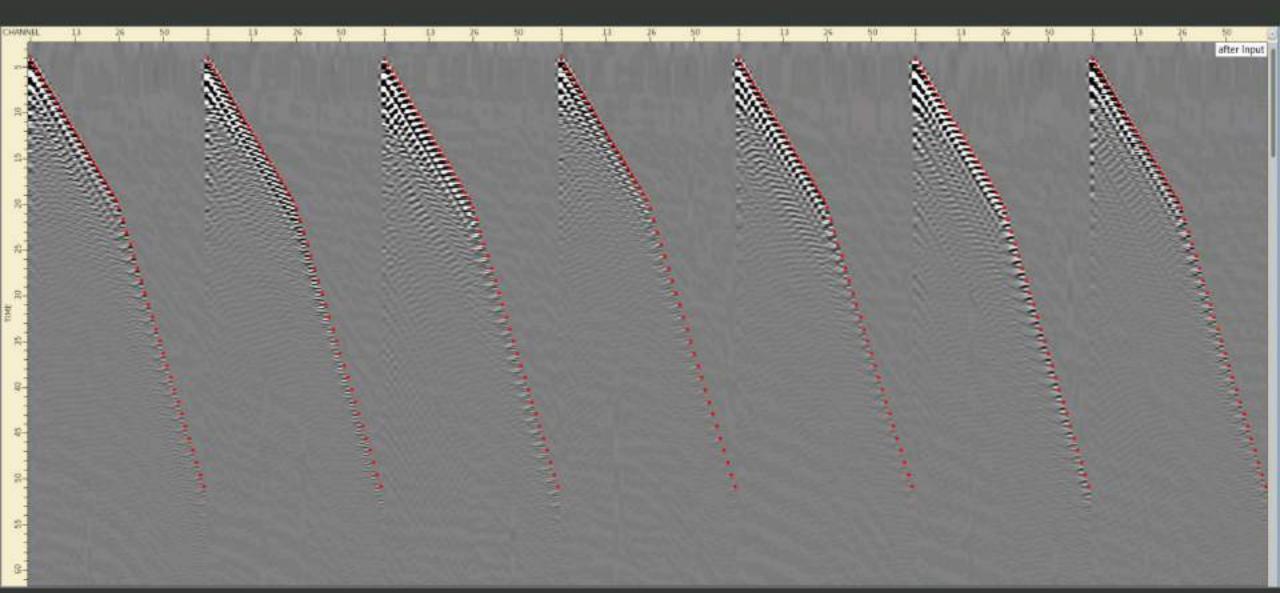


INCREASE SAMPLE RATE TO 0.1MS



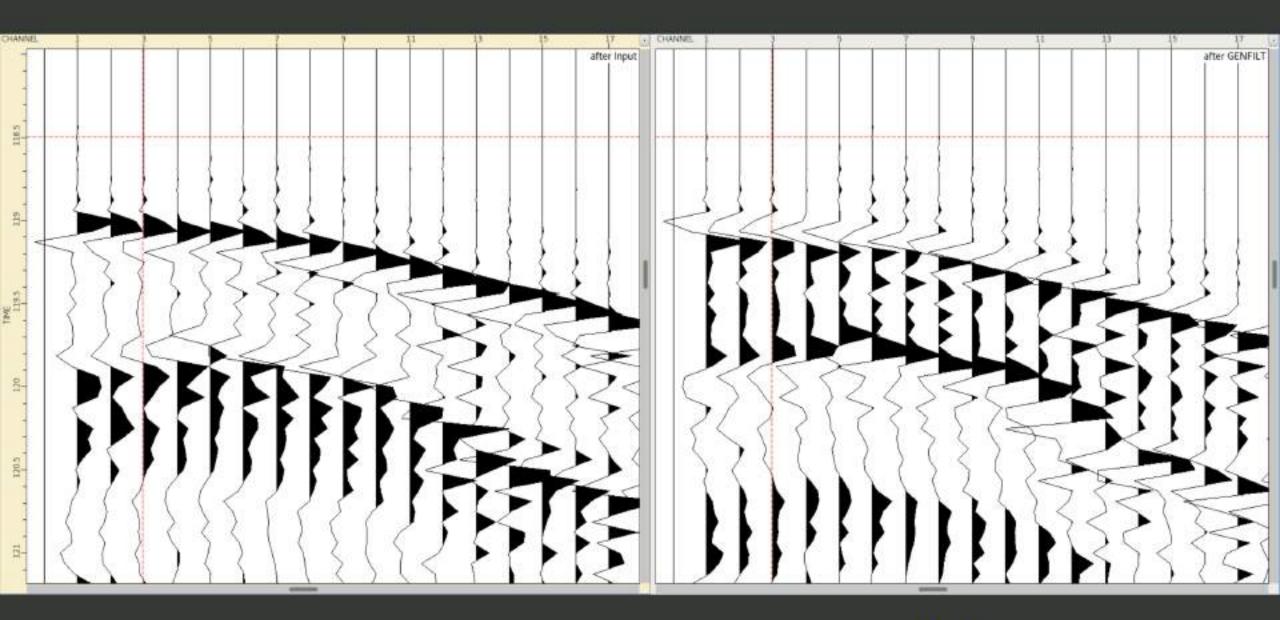


DIRECT ARRIVAL GEOMETRY QC





FLIP POLARITY



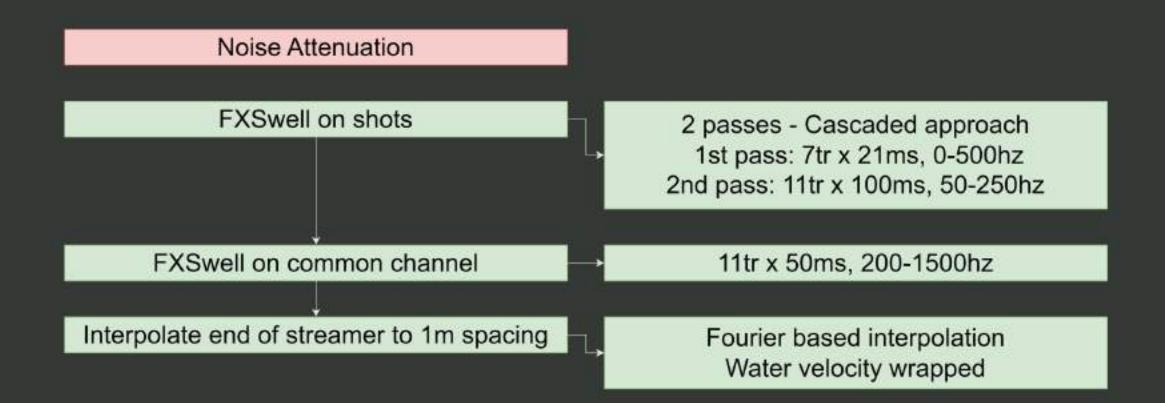


BRUTE STACK



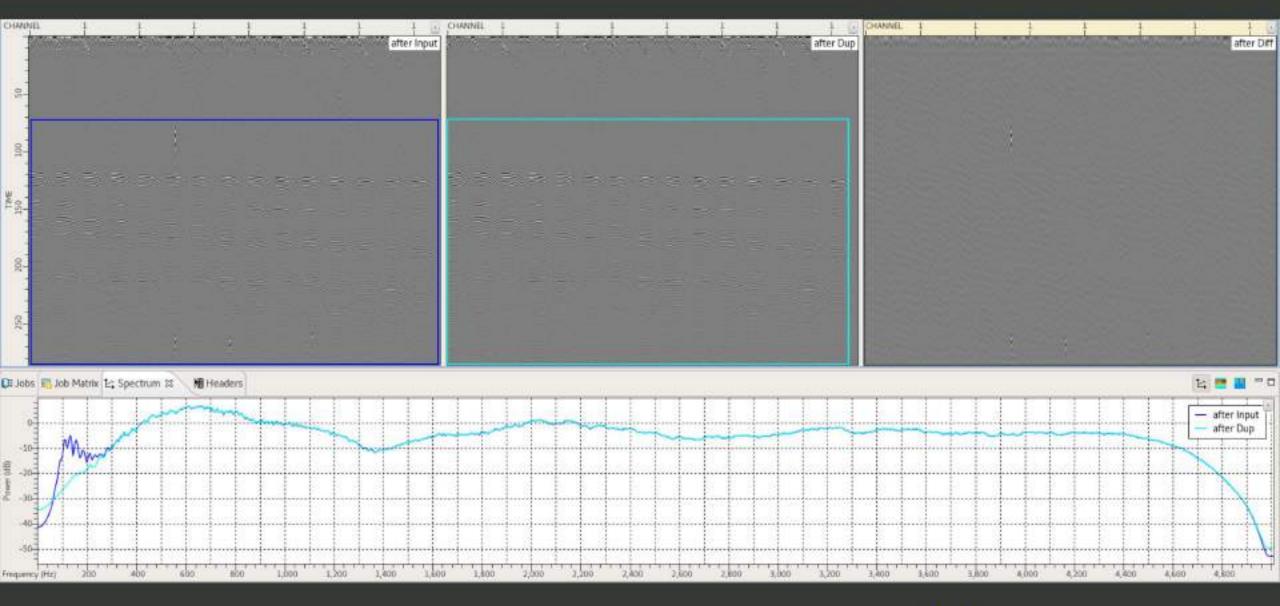


NOISE ATTENUATION



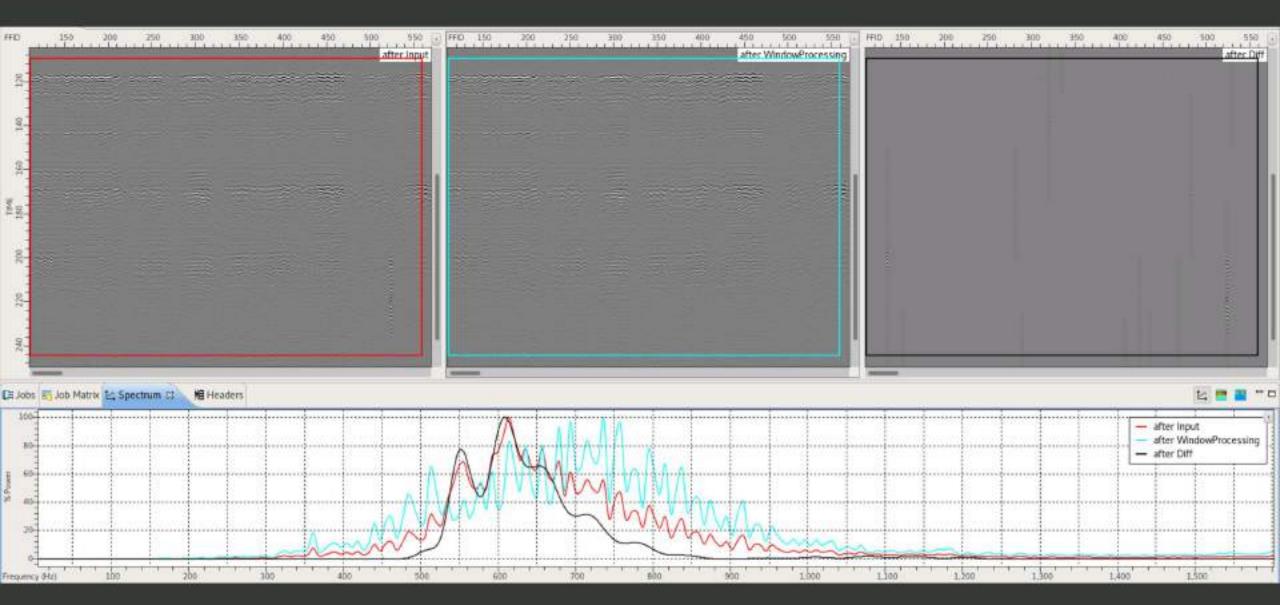


1ST PASS NOISE ATTENUATION – FXSWELL ON SHOTS



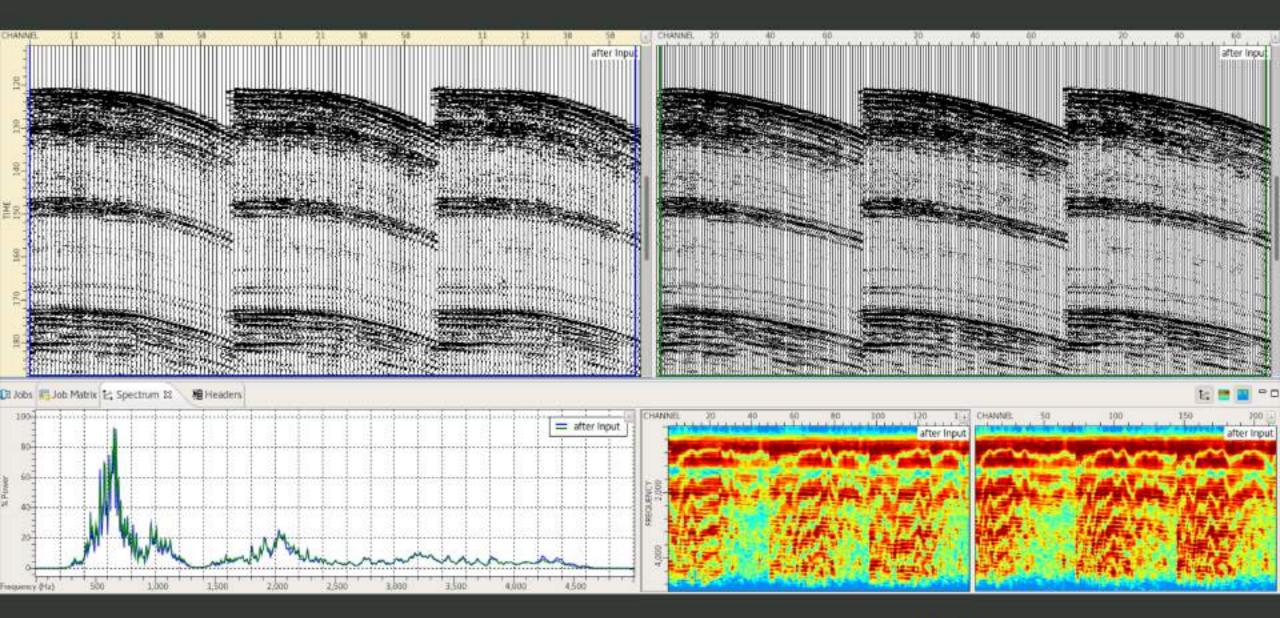


2ND PASS NOISE ATTENUATION – FXSWELL ON CHANNELS



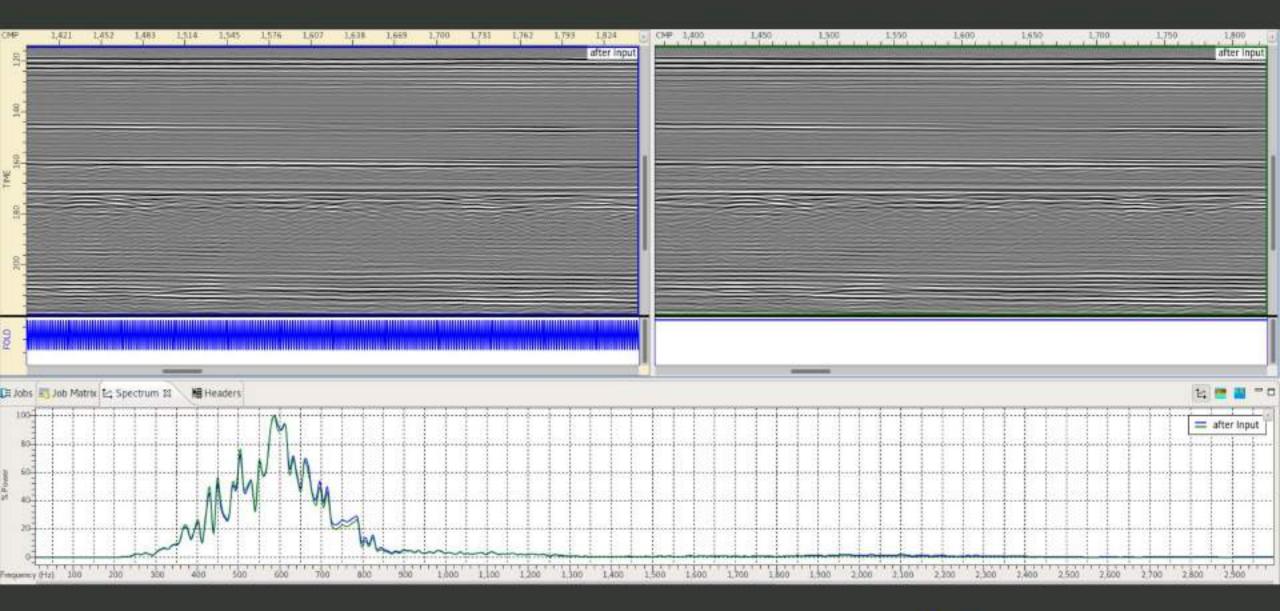


INTERPOLATE SHOTS TO 1M RECEIVER SPACING



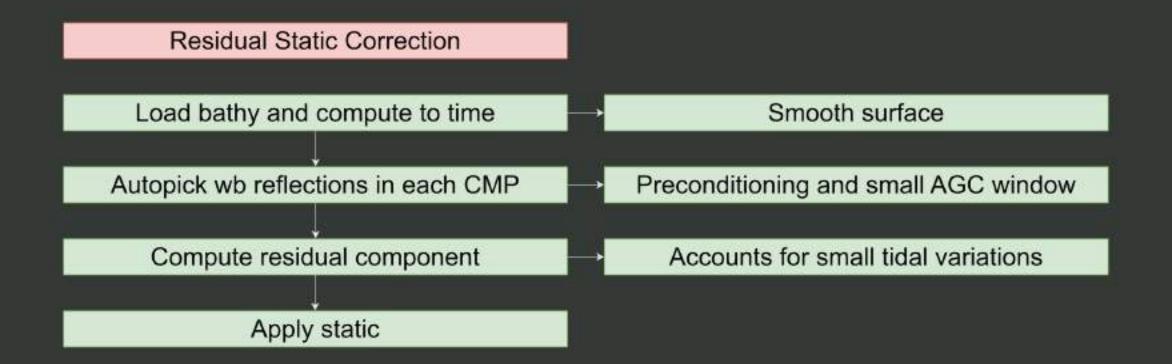


DENOISE STACK VS. INTERPOLATED STACK



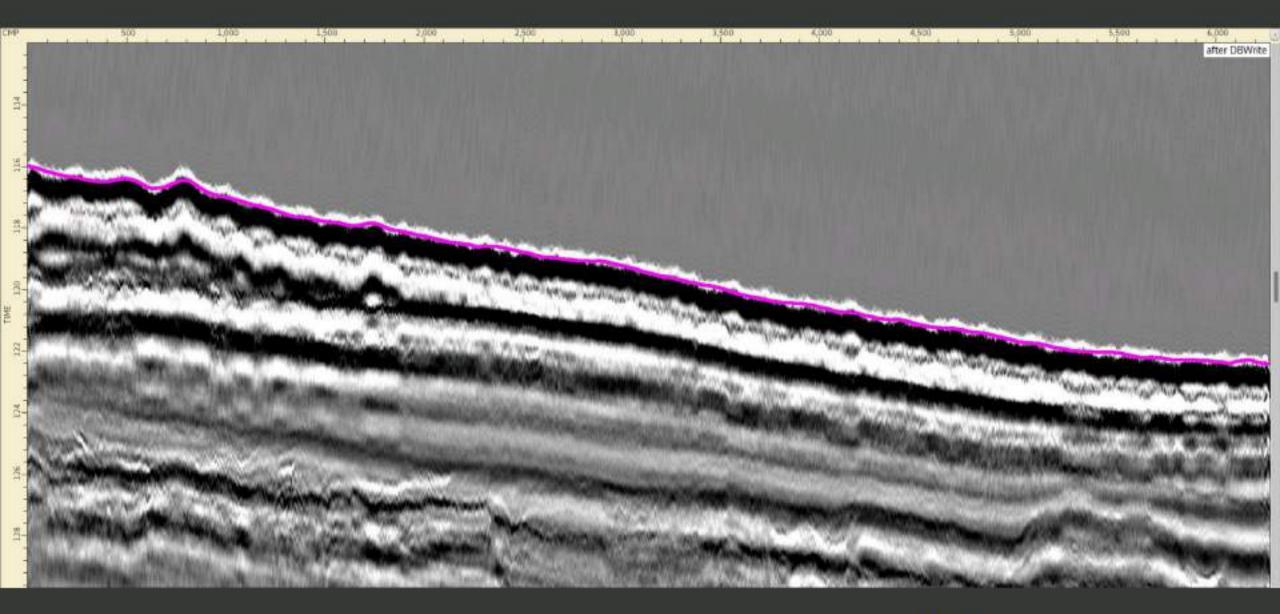


RESIDUAL STATIC CORRECTION



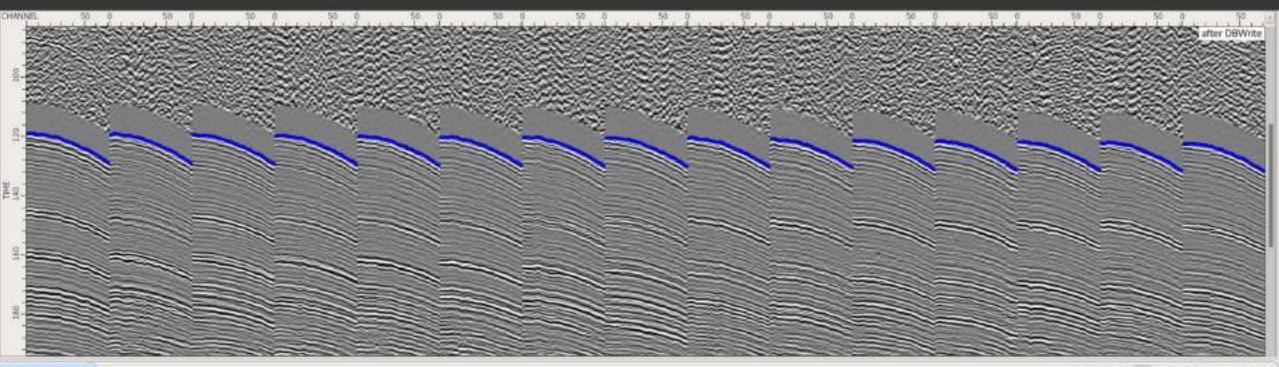


MULTIBEAM OVERLAY ON STACK





AUTOPICKED WB REFLECTION TIMES



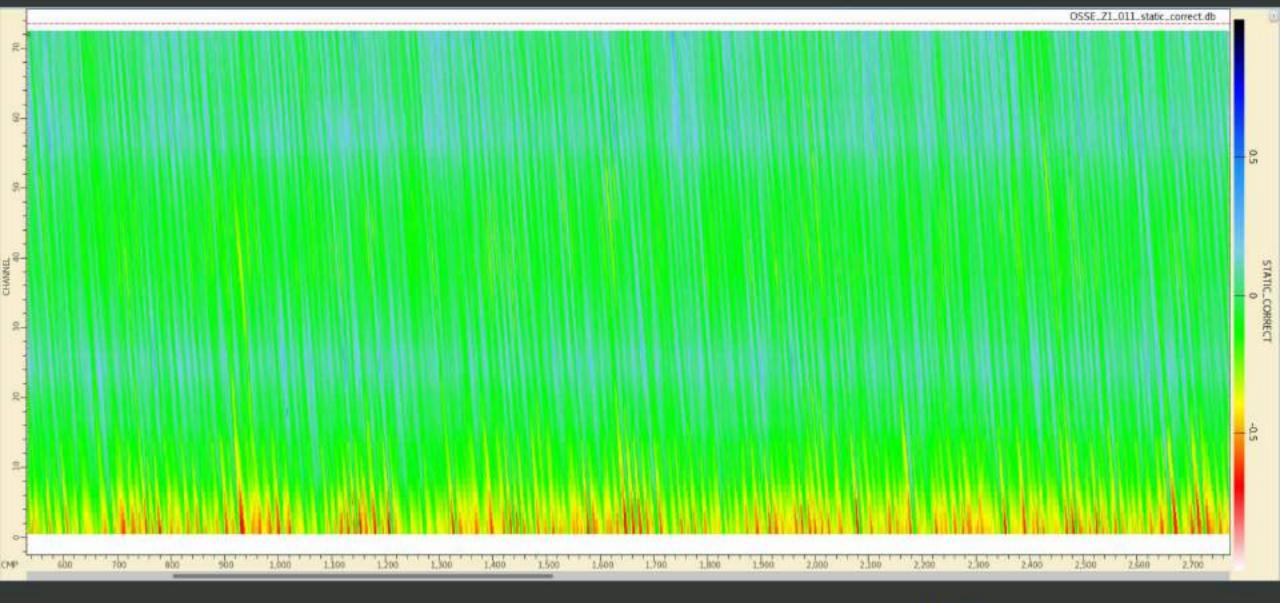
🛱 Data Navigation 🕄

1 - Q N N S K N Y -0

CMP	500 1,000	1,500	2,000 2,5	500 3,000	3,500 4,000	4,500 5,000 5,500 6,000
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				Second States Street Street		

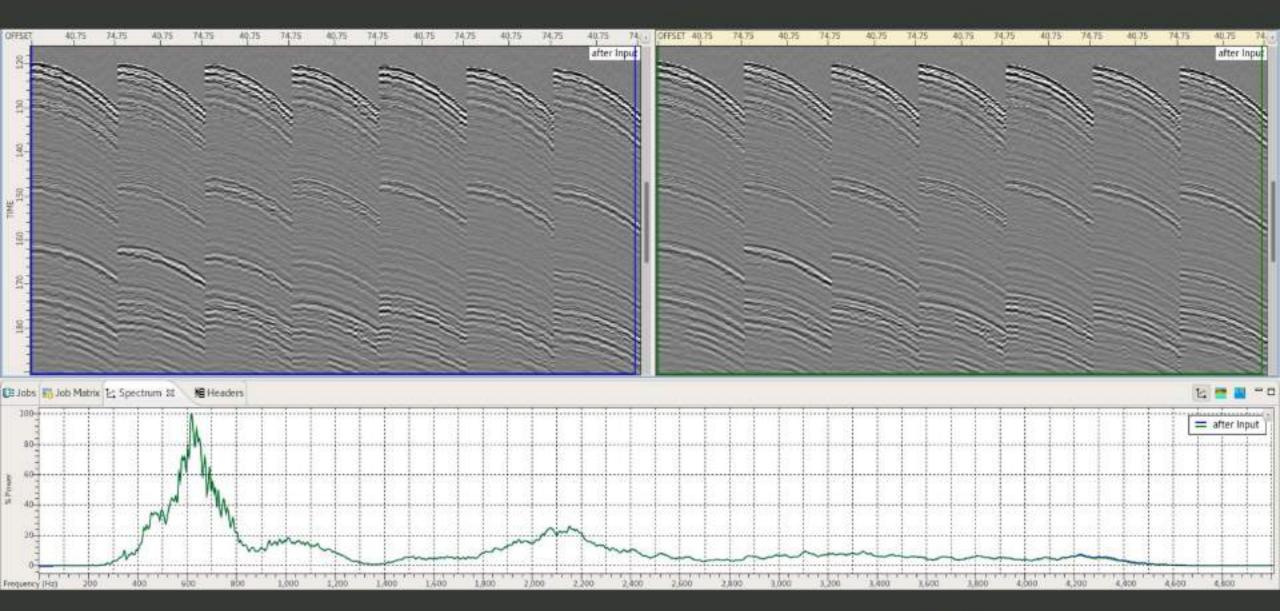


STATIC MAP



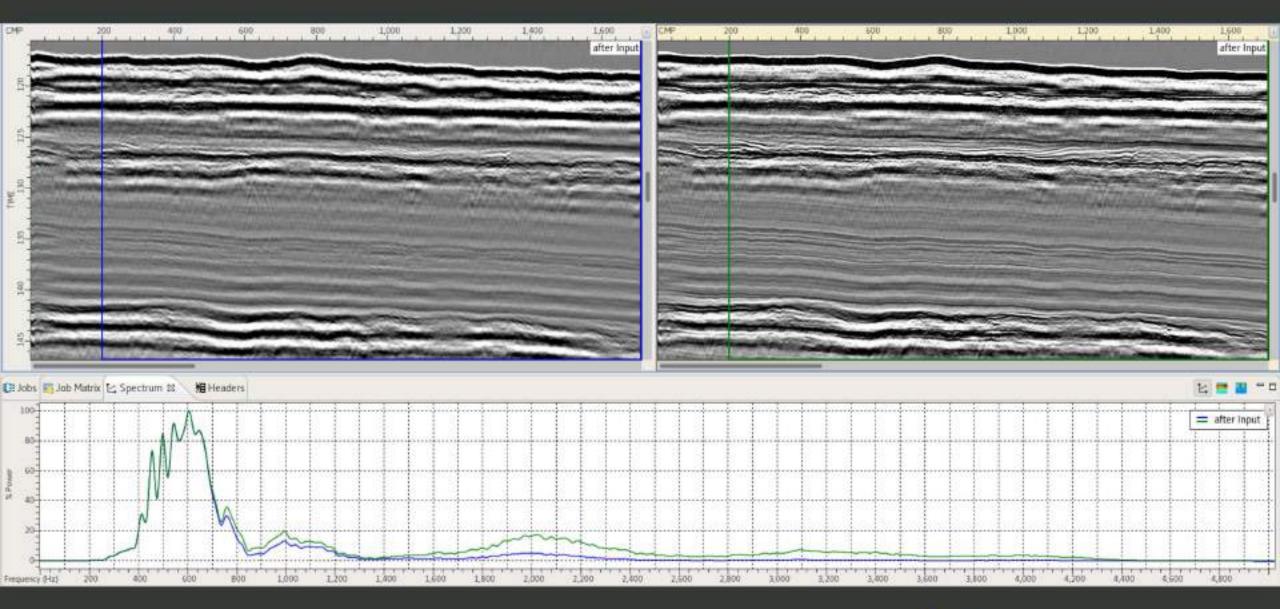


CMPS BEFORE AND AFTER CORRECTION



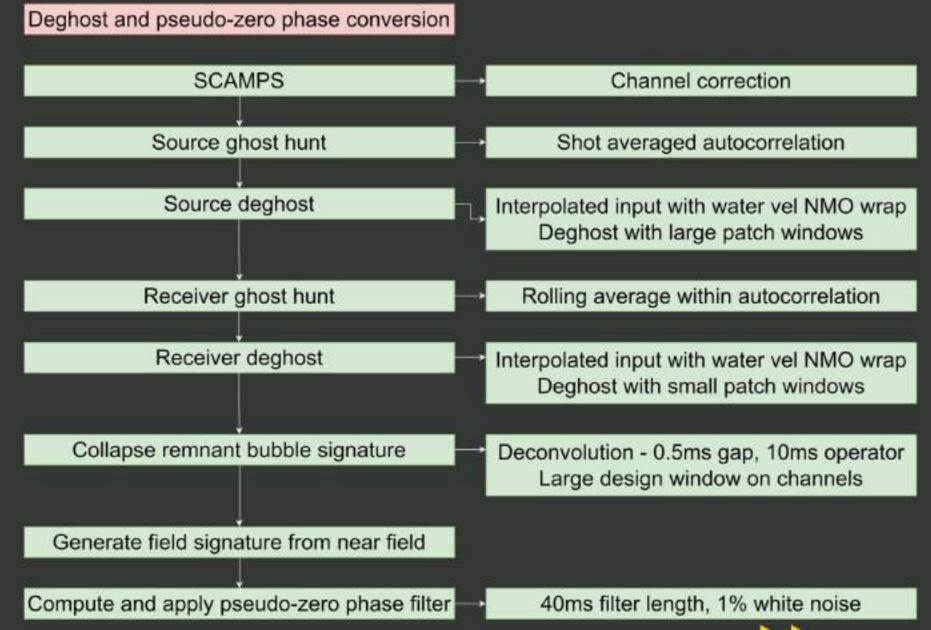


INTERPOLATED STACK VS. CORRECTED STACK



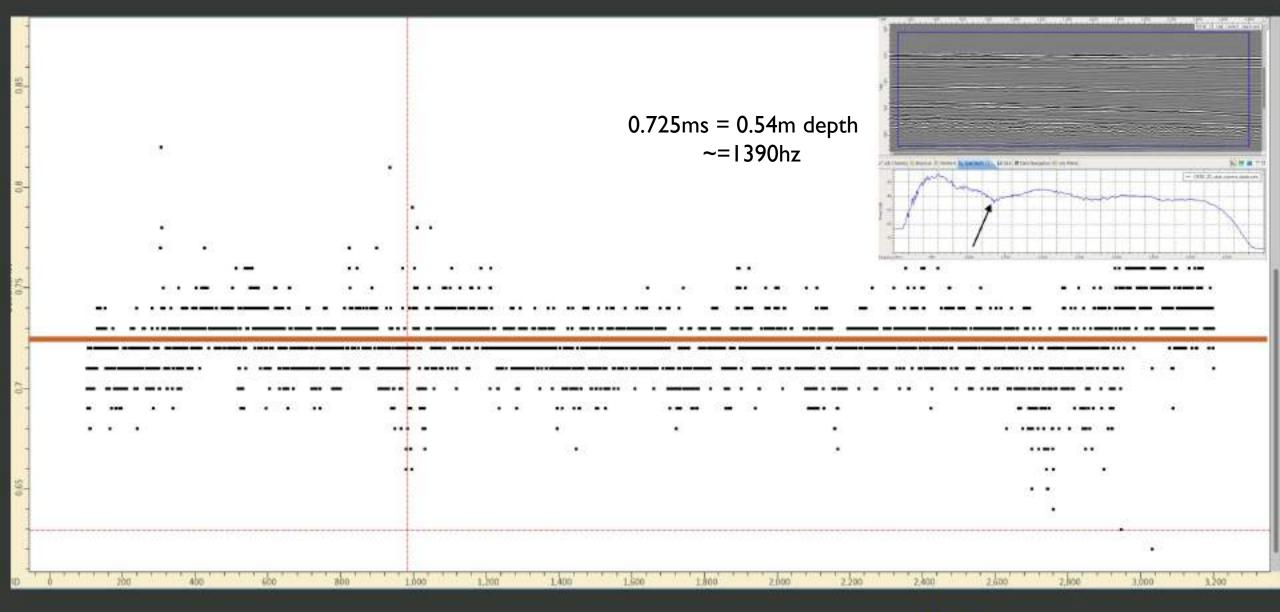


DEGHOST AND PSEUDO-ZERO PHASE CONVERSION



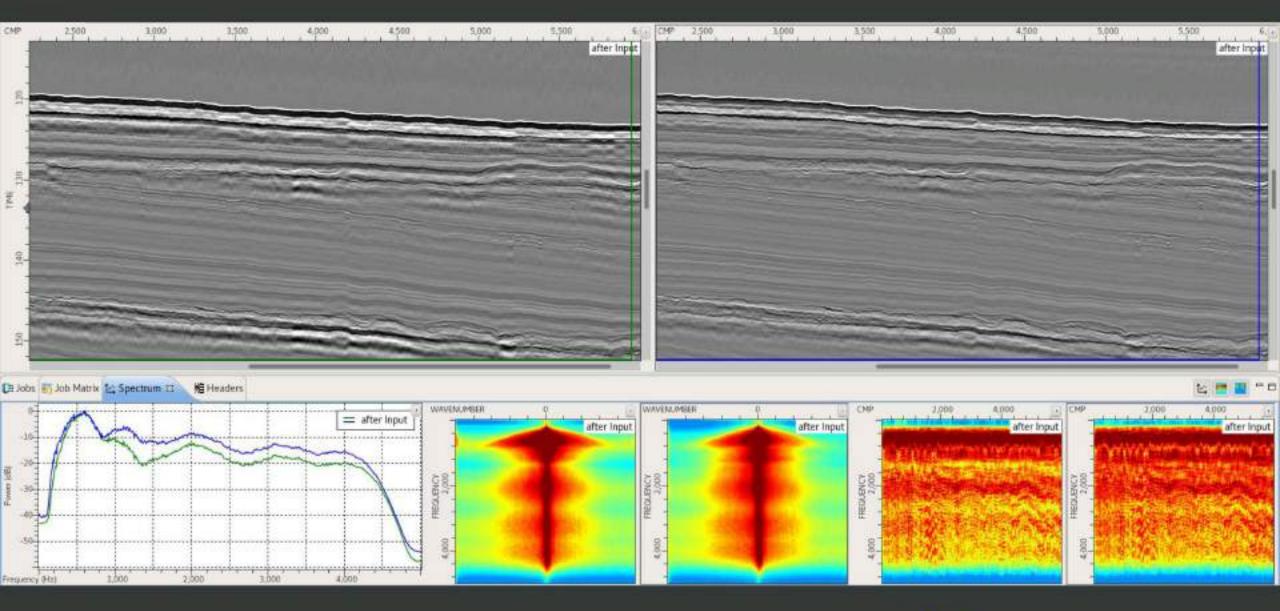


SOURCE DEPTH CALCULATION



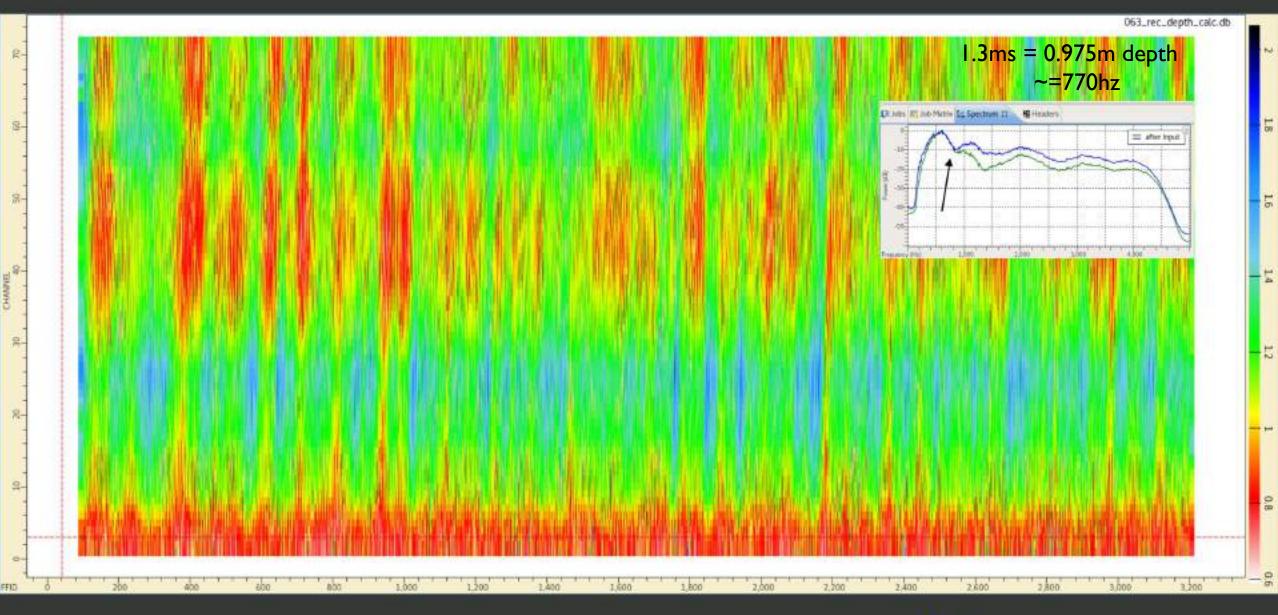


BEFORE AND AFTER SOURCE DEGHOSTING



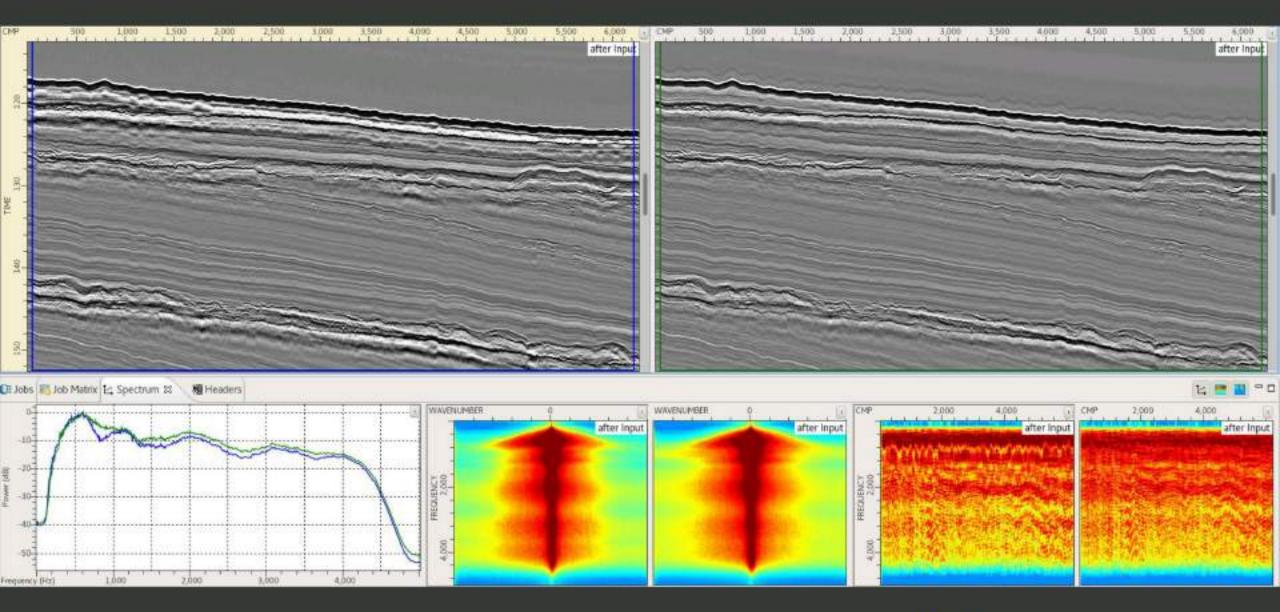


RECEIVER DEPTHS – FFID VS. CHANNEL



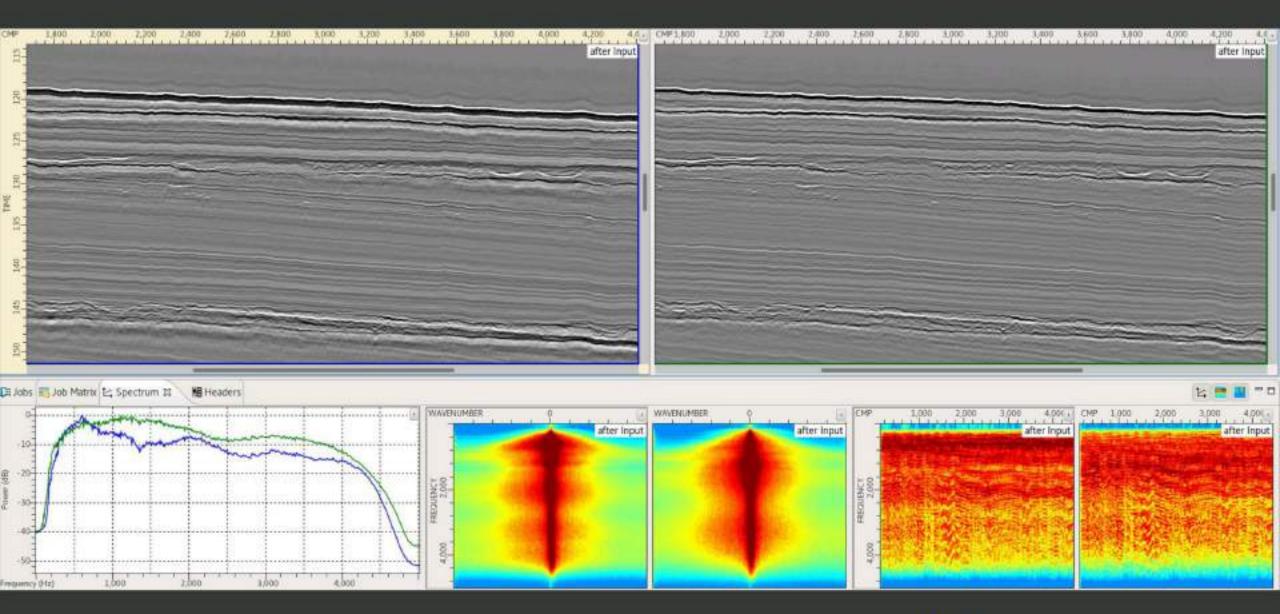


BEFORE AND AFTER RECEIVER DEGHOSTING



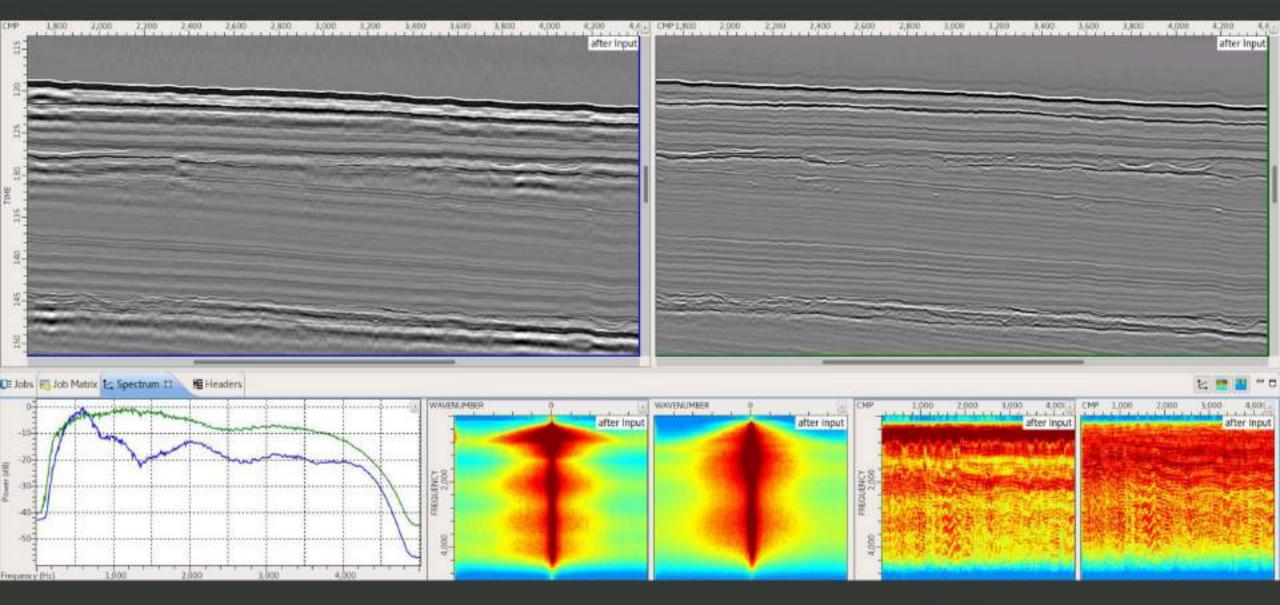


BEFORE AND AFTER WAVELET COMPRESSION



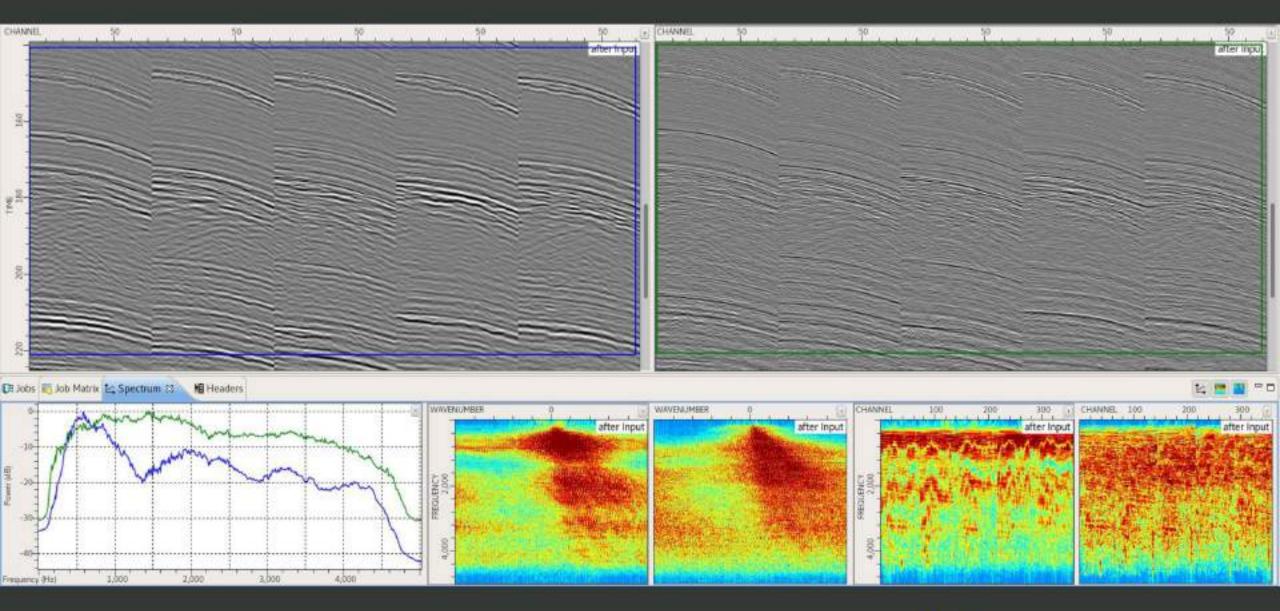


BEFORE AND AFTER DEGHOSTING



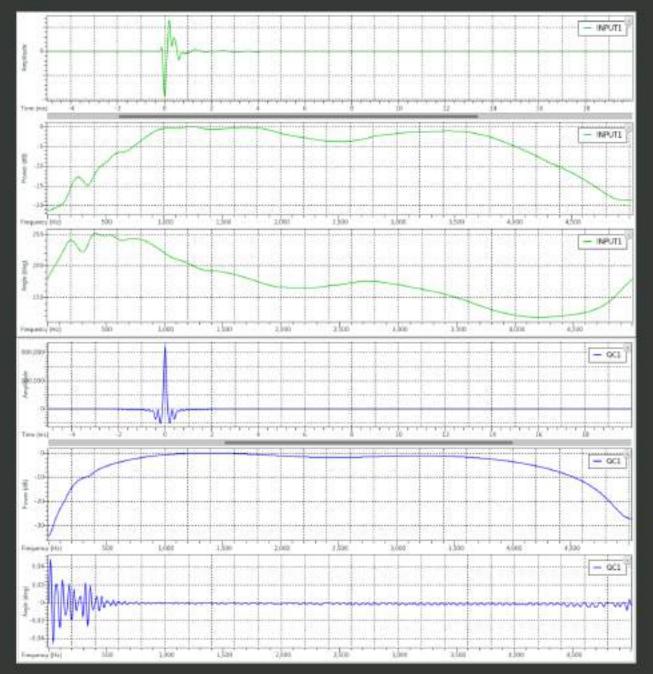


SHOTS BEFORE AND AFTER DEGHOSTING



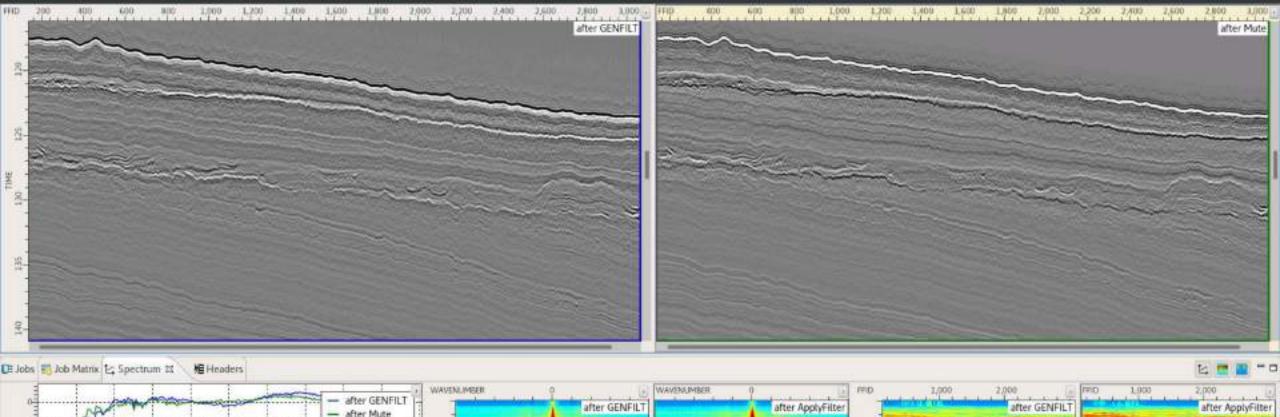


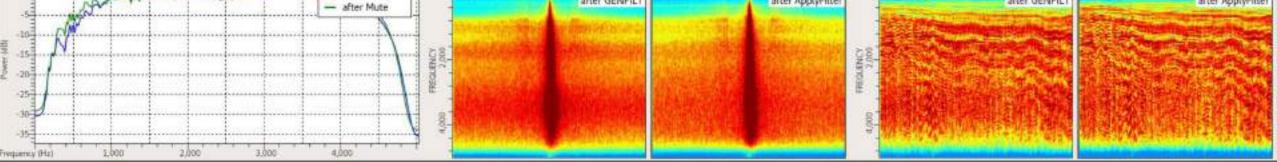
WAVELET DESIGN





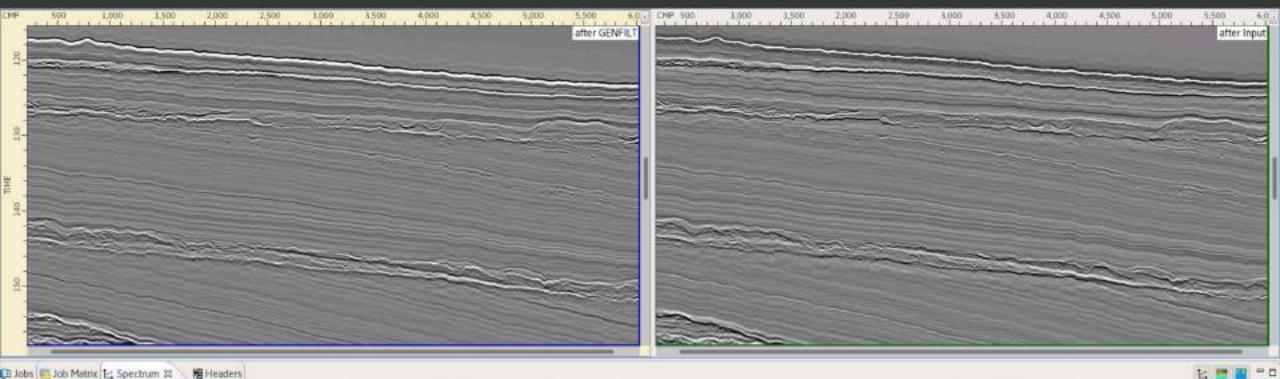
CHANNEL 10 – BEFORE AND AFTER PSEUDO-ZERO PHASE

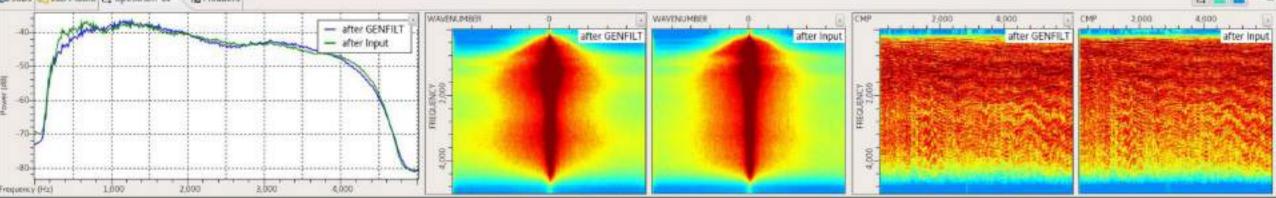






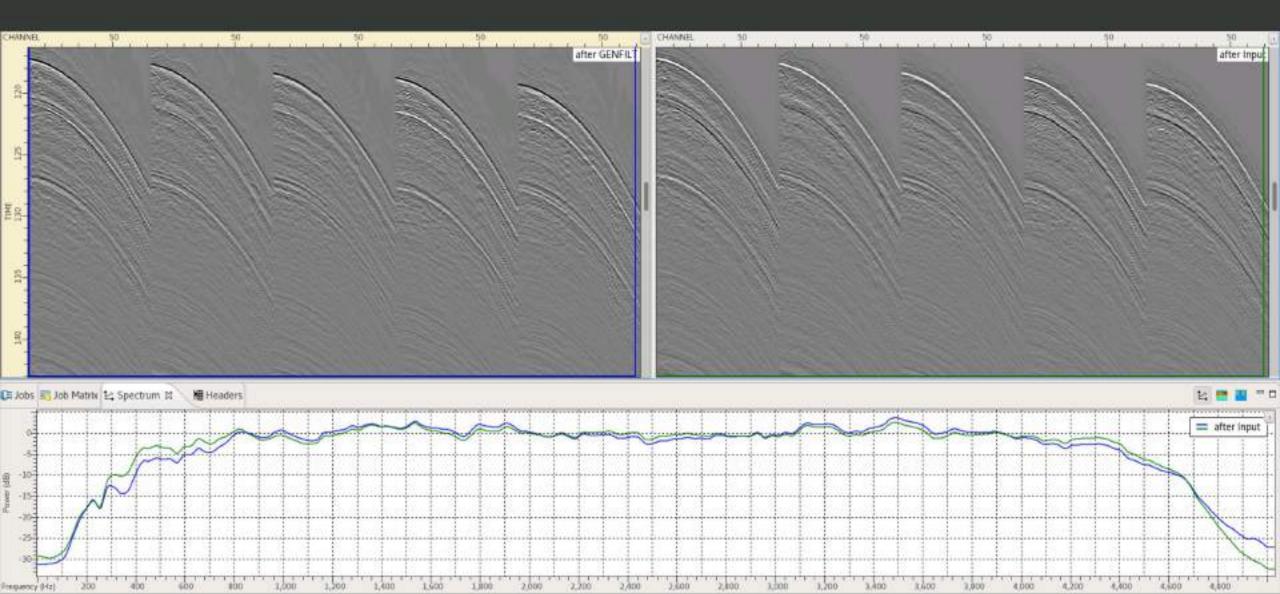
STACKS BEFORE AND AFTER PSEUDO-ZERO PHASING





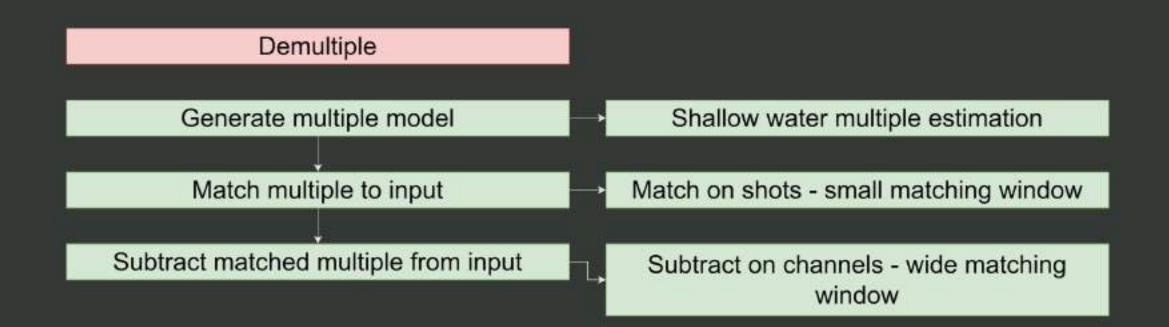


SHOTS BEFORE AND AFTER PSEUDO-ZERO PHASING



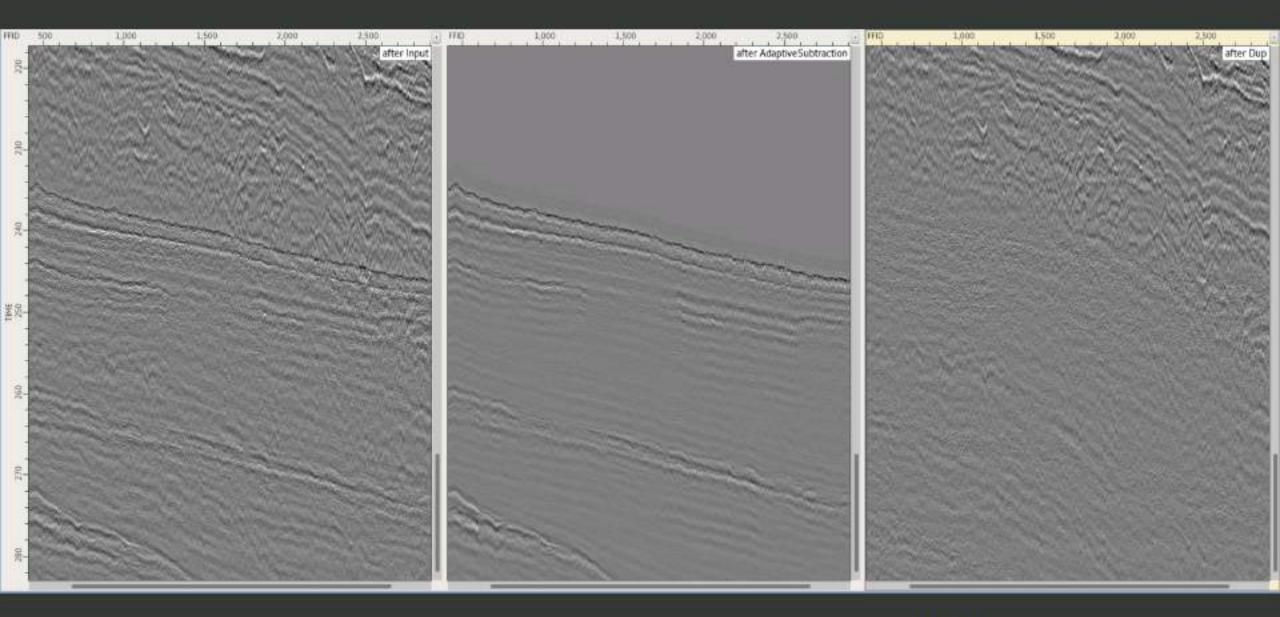


DEMULTIPLE



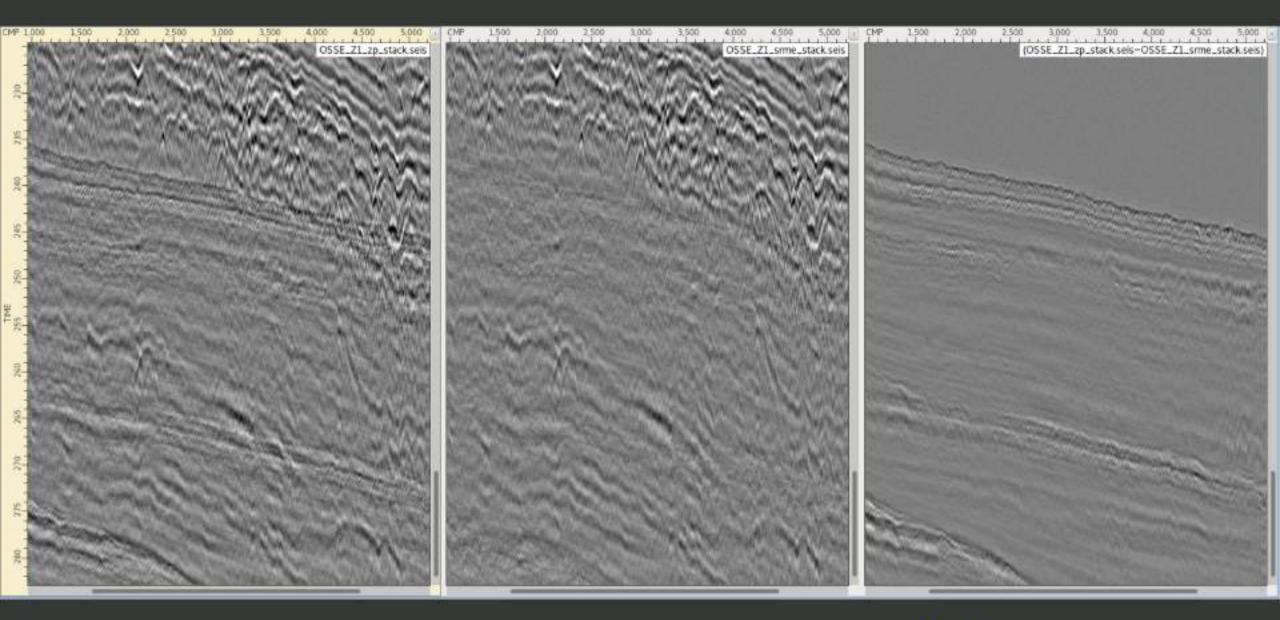


CHANNEL 1 – BEFORE, MODEL AND AFTER SUBTRACTION



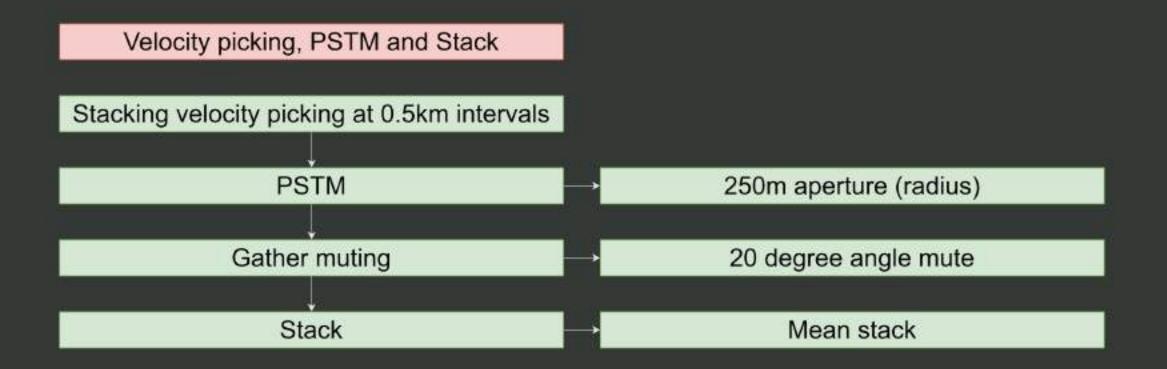


STACKS – BEFORE, AFTER AND DIFFERENCE



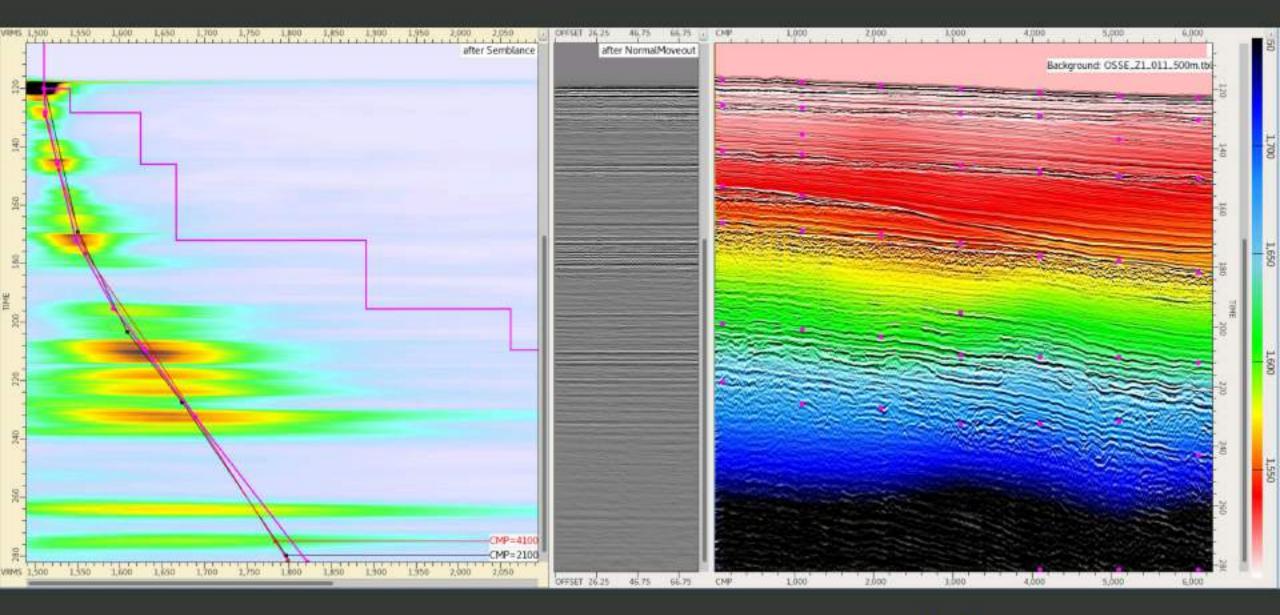


VELOCITY PICKING, PSTM AND STACK



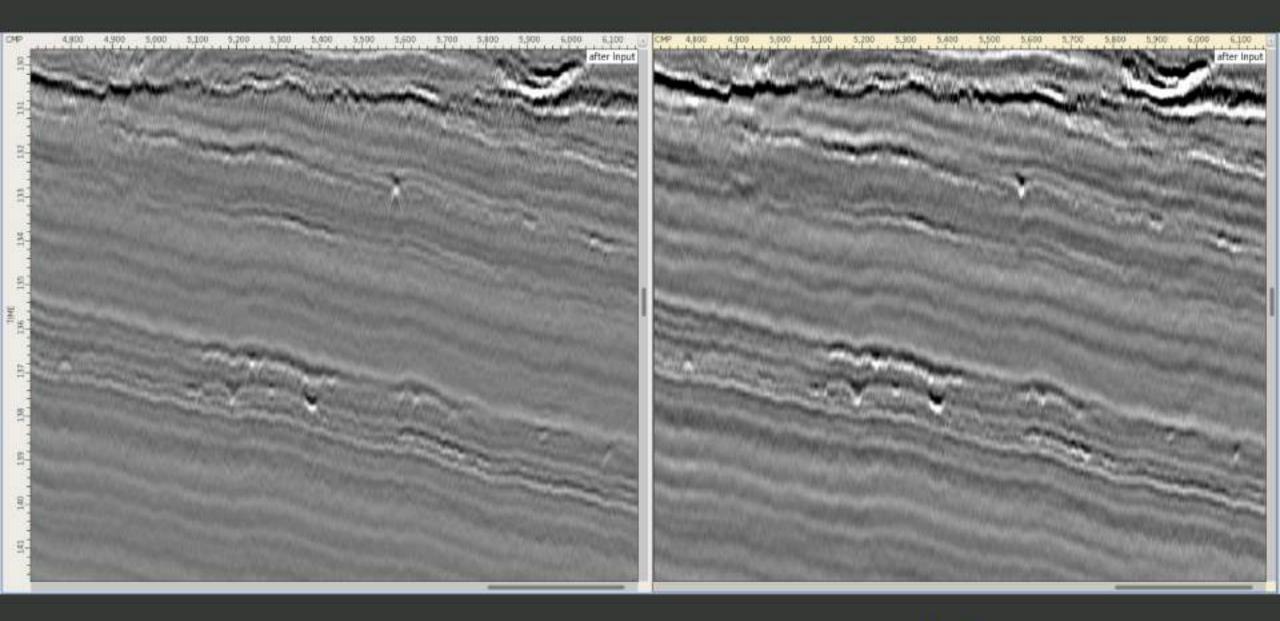


VELOCITY PICK EXAMPLE AT 0.5KM INTERVAL



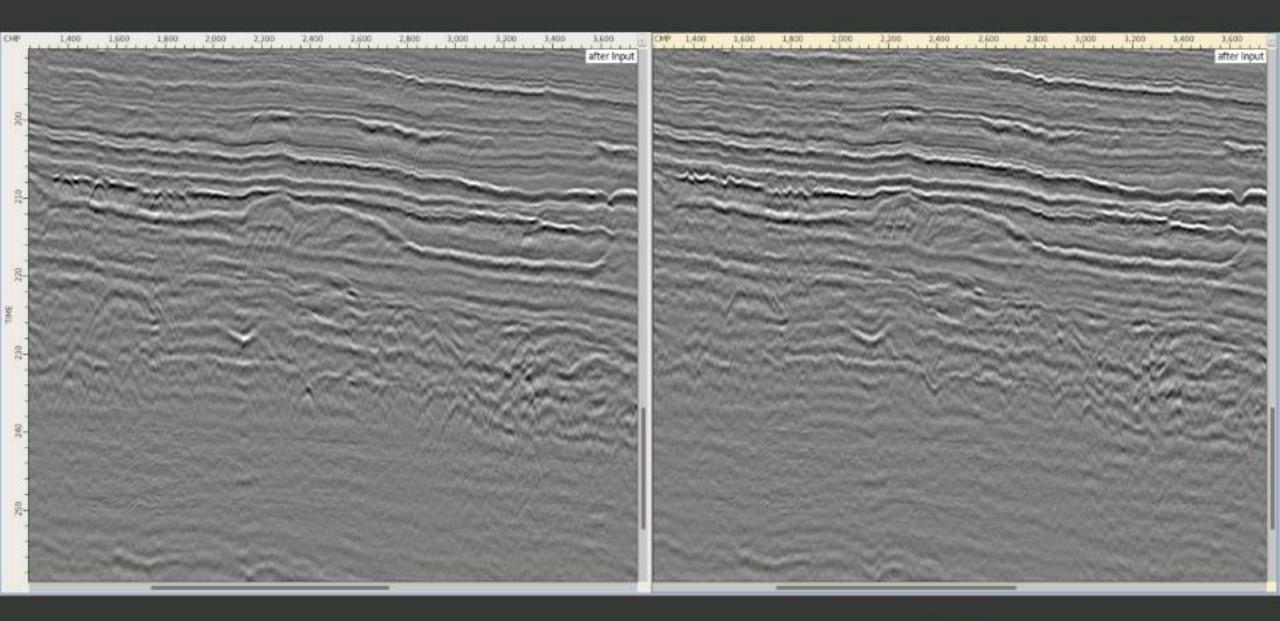


STACKS BEFORE, AFTER MIGRATION – SHALLOW SECTION



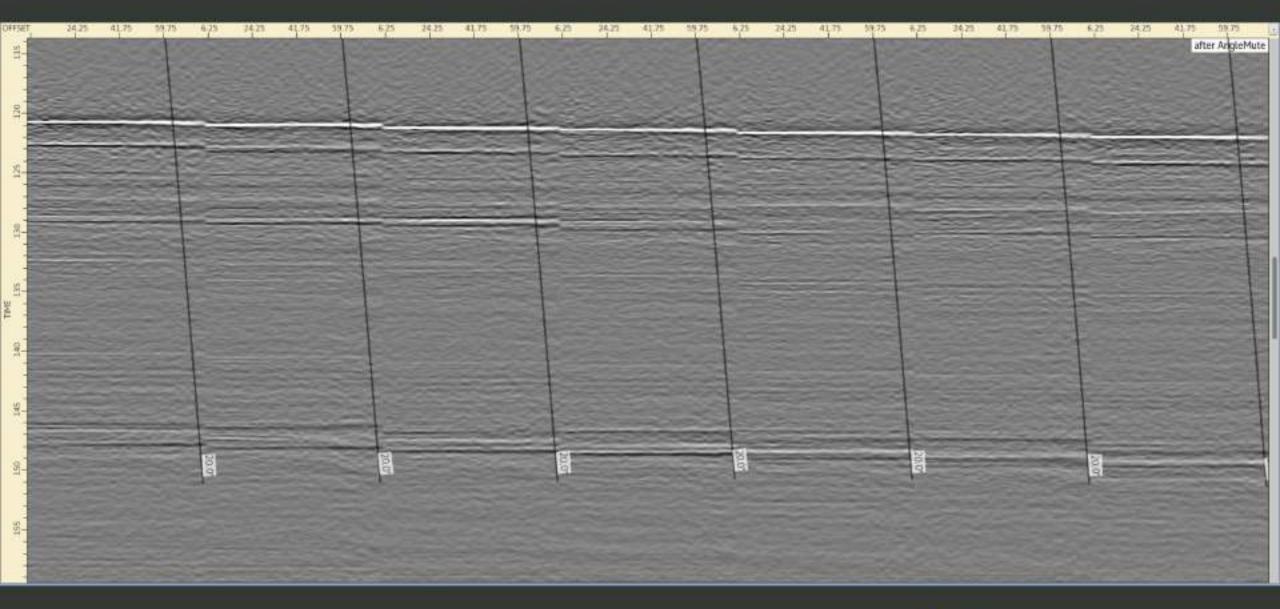


STACKS BEFORE, AFTER MIGRATION – DEEPER SECTION



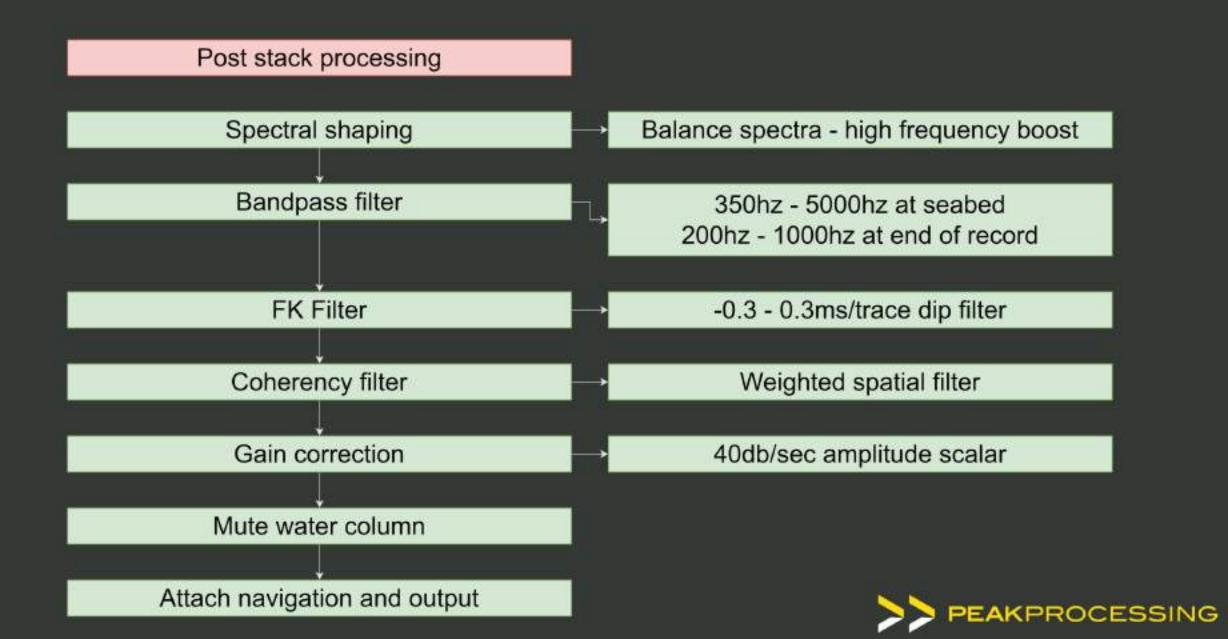


GATHERS – PROPOSED ANGLE MUTE

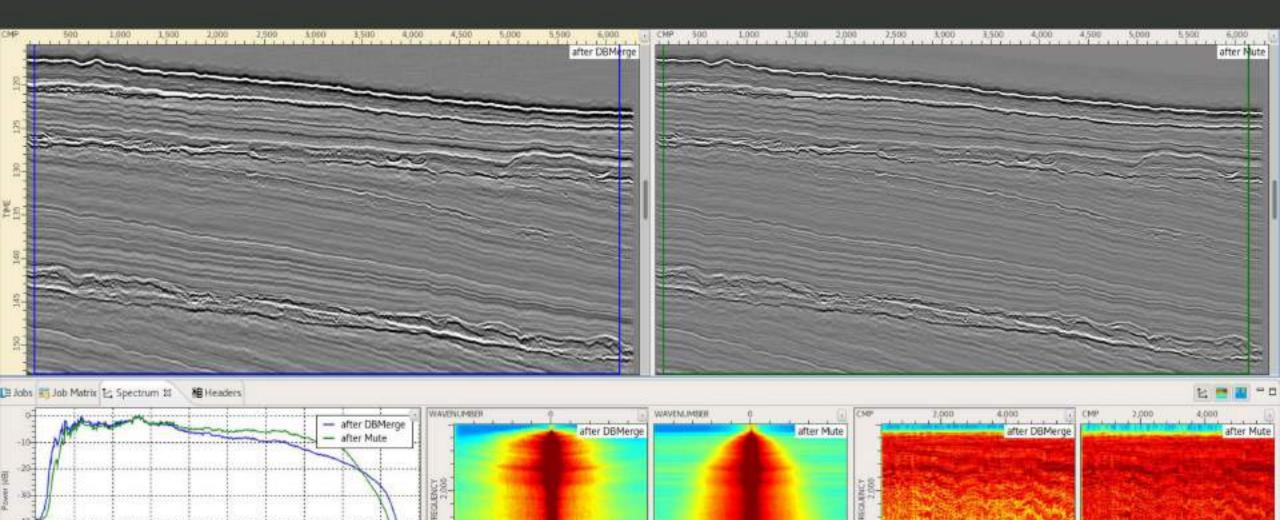




POST STACK PROCESSING



BEFORE AND AFTER POST STACK PROCESSING



4,000

Frequency (Hz)

1,000

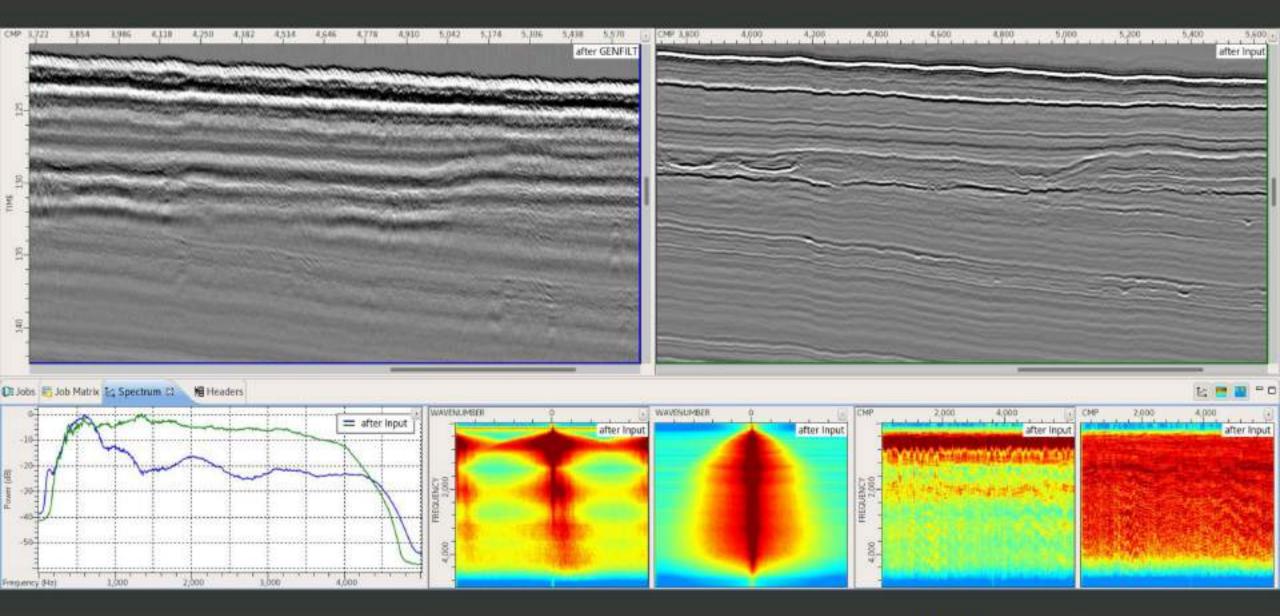
2,000

3,000

4,000

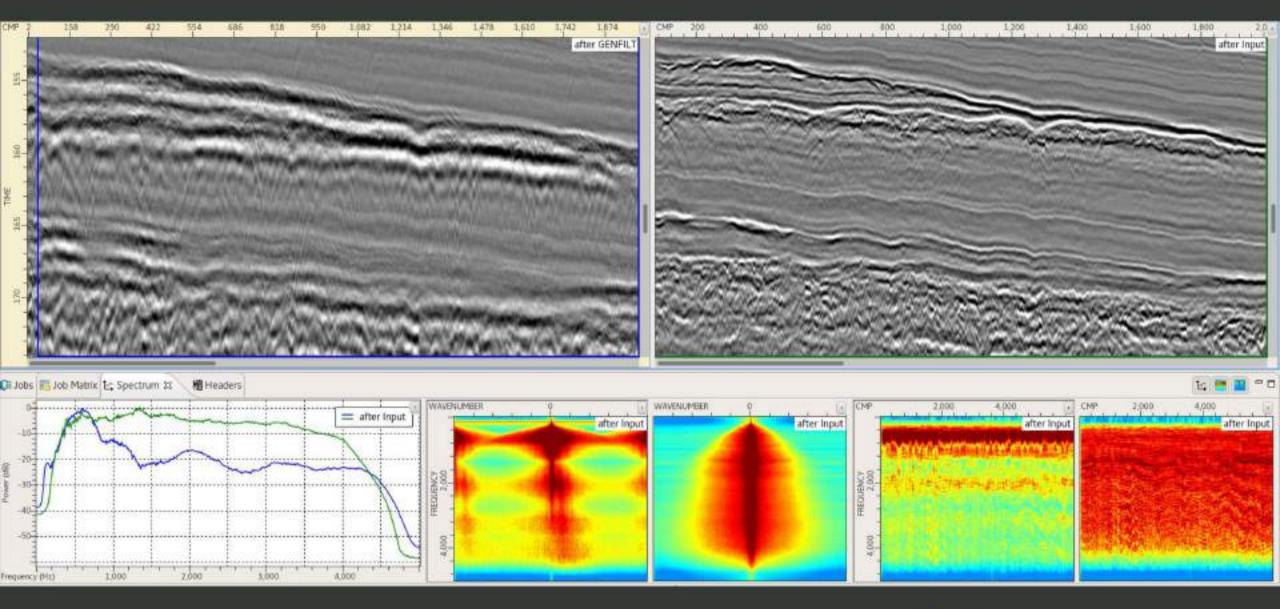


BRUTE VS. FINAL COMPARISON



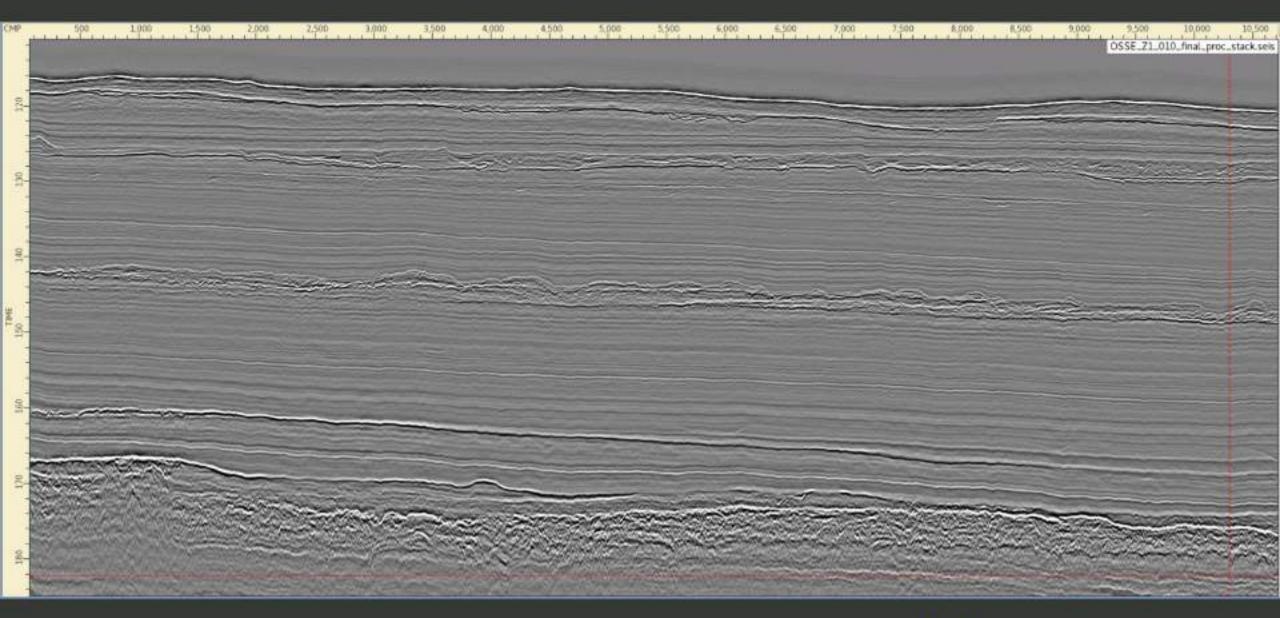


BRUTE VS. FINAL COMPARISON





OSSE_Z1_010 FINAL EXAMPLE





OSSE_Z1_017 FINAL EXAMPLE

